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Nonlinear ice sheet/ liquid interaction in a channel with an obstruction

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The interaction between the flow in a channel with an obstruction on the bottom and an elastic sheet representing the ice covering the liquid is considered for the case of steady flow. The mathematical model based on the velocity potential theory and the theory of thin elastic shells fully accounts for the nonlinear boundary conditions at the elastic sheet/liquid interface and on the bottom of the channel. The integral hodograph method is employed to derive the complex velocity potential of the flow, which contains the velocity magnitude at the interface in explicit form. This allows one to formulate the coupled ice/liquid interaction problem and reduce it to a system of nonlinear equations in the unknown magnitude of the velocity at the interface. Case studies are carried out for a semi-circular obstruction on the bottom of the channel. Three flow regimes are studied: a subcritical regime, for which the interface deflection decays upstream and downstream; an ice supercritical and channel subcritical regime, for which two waves of different lengths may exist; and a channel supercritical regime, for which the elastic wave is found to extend downstream to infinity. All these regimes are in full agreement with the dispersion equation. The obtained results demonstrate a strongly nonlinear interaction between the elastic and the gravity wave near the first critical Froude number where their lengths approach each other. Results for the interface shape, the bending moment, and the pressure along the interface are presented for wide ranges of the Froude number and the obstruction height.

1. Introduction

The problem of the interaction between a liquid and an elastic boundary is a classical problem in fluid mechanics, which has applications in offshore and polar engineering, medicine, and various industrial fields. In recent decades, this topic has gained renewed attention due to global warming and the melting of ice in Arctic regions, which has opened up new routes for ships and new areas for resource exploration (Squire et al. (1995), Părău and Dias (2002), Korobkin, Părău and Vanden-Broeck (2011), Blyth, Părău and Vanden-Broeck (2011)).

In the past century, studies on ice/liquid interaction primarily focused on the response of an ice cover to a load moving on the ice surface. This problem was driven by the practical need for seasonal routes for vehicles and runways for aircraft in polar regions (Squire et

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36 al. (1988)). A comprehensive bibliography on this subject can be found in the monograph
37 Squire et al. (1996).

38 Current studies on ice-related phenomena are centered around the effect of ice on ocean
39 waves and their interaction with various ice structures, such as continuous ice, floes, polynyas,
40 and pancake ice. One important aspect is understanding how far ocean waves can penetrate
41 into ice fields, leading to the breaking of ice near the shore and the formation of a marginal
42 ice zone with multiple cracks and polynyas (Guyenne and Părău (2012), Guyenne and Părău
43 (2017), Meylan et al. (2018), Squire (2020)).

44 Studying the interaction between an ice sheet and water waves is mathematically chal-
45 lenging. Most publications in this field rely on linear theories of water waves and the theory
46 of a thin elastic shell to model the ice cover (Sturova (2009), Karmakar (2010), Korobkin,
47 Părău and Vanden-Broeck (2011), Khabakhpasheva et al. (2019), Shishmarev et al. (2019),
48 Stepanyants and Sturova (2021)). One interesting aspect of ice/water interaction is different
49 types of ice response depending on the wave velocity caused by a moving disturbance, such
50 as a load on the ice sheet or a body moving beneath the ice sheet. Linear theories can be
51 used to derive the dispersion relation and determine two critical wave speeds: one applies
52 to gravity waves in a channel of finite depth, and the other is the minimal speed of wave
53 propagation at the interface due to the elastic sheet (Kheisin (1963), Kheisin (1967)). [The](#)
54 [corresponding critical Froude numbers based on the depth of the channel are denoted as](#)
55 [\$F = 1\$ and \$F = F_{cr}\$.](#)

56 For wave speeds in the range between these two critical speeds, the linear theories predict
57 two waves of different lengths: a longer wave due to gravity moving downstream, and a shorter
58 wave moving upstream caused by the elastic sheet. A linear theory is also employed to study
59 ice/water/structure interaction, with recent reviews provided by Ni et al. (2020). Some papers
60 in this field focus on the effects of bottom topography and an arbitrary ice thickness. For
61 example, Porter and Porter (2004) used a variational approach to study the effect of varying
62 the ice thickness and the water depth on wave propagation in three dimensions. Sturova
63 (2009) investigated the unsteady behavior of ice floating on shallow water with a variable
64 depth. Karmakar (2010) analyzed wave transformation by multiple steps and blocks on the
65 channel bottom using the wide-spacing approximation. Shishmarev et al. (2019) explored
66 methods to mitigate oscillations of floating elastic plates under periodic surface water waves.
67 Ice response on load moving along on frozen channel and on motion of underwater body
68 was investigated by Shishmarev et al. (2016), Shishmarev et al. (2019) and Shishmarev et al.
69 et al. (2023). At the last work thickness of ice cover was variable across a channel. Large time
70 response of ice cover on underwater moving body was described in Khabakhpasheva et al.
71 (2019). Xue et al. (2021) investigated the hydroelastic response of an ice sheet with a lead
72 to a moving load.

73 However, the linear theories cannot accurately predict the behavior of an ice sheet near the
74 critical speed, where they predict an infinite response of the interface. Nonlinear studies of
75 flexural-gravity waves in this context are limited. Părău and Dias (2002) studied the effects
76 of nonlinearity slightly below the critical wave [speed, or \$F < F_{cr}\$](#) , and derived a nonlinear
77 Schrödinger equation. Bonnefoy et al. (2009) developed a higher-order spectral method to
78 calculate the nonlinear response of an infinite ice sheet to a moving load in the time domain.
79 Milewski et al. (2011) obtained purely hydroelastic solitary waves for a full nonlinear
80 model in deep water using a conformal mapping technique. Gao et al. (2019)) extended this
81 method to finite depth flows with constant vorticity. Guyenne and Părău (2012) discovered
82 depression and elevation branches of solitary waves below the minimum phase speed using
83 the Cosserat theory of hyperelastic shells satisfying Kirchhoff's hypotheses (Plotnikov and
84 Toland (2011)). They compared the wave profiles computed by the boundary-integral method

85 and high-order spectral method. Strongly nonlinear events were also studied for a jet impact
86 on an ice sheet (Yuan et al. (2022)) and for ice–bubble interaction (Zhang et al. (2023)).

87 The nonlinear studies mentioned above mainly focus on exploring solitary waves with an
88 ice sheet in deep or constant depth water. Both the steady and the unsteady formulations
89 of the problem are used to predict the wave propagation originated by the pressure load on
90 the ice sheet. Page and Părău (2014) investigated the steady problem of hydraulic fall in
91 the presence of an ice sheet and bottom geometry. They used the Cosserat theory to model
92 the ice sheet and employed boundary integral equation techniques to solve the problem for
93 the liquid region. They presented results for hydraulic falls without wave trains upstream or
94 downstream; however, they obtained solutions with a train of waves trapped between two
95 obstructions.

96 In this paper, a general solution to the steady nonlinear problem of hydroelastic waves
97 generated by an obstruction on the channel bottom is presented. The problem is equivalent to
98 a body moving beneath an ice sheet along a flat bottom in still water. Although the formulation
99 of the problem is steady and two-dimensional, that is, simpler than the unsteady formulations
100 in the studies mentioned above, the present study focuses on the nonlinear features of the
101 elastic sheet /fluid interaction which have not been explored before. In particular, how
102 the height of the obstruction affects the interface, the bending moment, and the pressure
103 distribution along the interface in the whole range of flow velocities, including the subcritical
104 and the supercritical flow regime; at what maximal height of the obstruction a steady solution
105 still exists. For supercritical flows with Froude number $F > 1$, the present study revealed the
106 existence of flexural gravity waves downstream of the obstruction, which are in agreement
107 with those predicted by the dispersion relation. The integral hodograph method is employed to
108 derive the complex velocity potential, which includes the velocity magnitude at the ice/liquid
109 interface and the slope of the bottom in explicit form. The coupling of the elastic sheet and
110 moving liquid solutions is based on the condition of an equal pressure at the interface, which
111 arises both from flow dynamics and from elastic sheet equilibrium. The entire problem
112 is reduced to a system of nonlinear equations in the unknown velocity magnitude at the
113 interface, which is solved numerically. This methodology was previously applied to infinite
114 depth water (Semenov (2021)) and to the flow in a channel covered by broken ice (Ni et al.
115 (2023)).

116 The derivation of the flow potential and the numerical method for solving the coupled
117 liquid/elastic sheet interaction problem are presented in Section II. Extended numerical
118 results are discussed in Section III. The solution is carefully checked by reproducing the
119 results of Page and Părău (2014) for the hydraulic fall under an ice plate. Then, three
120 flow regimes are studied: a subcritical regime ($F < F_{cr}$), an ice supercritical and channel
121 subcritical regime ($F_{cr} < F < 1$), and a channel supercritical regime ($F > 1$). For the Froude
122 number range $F_{cr} < F < 1$, the presented results revealed a strongly nonlinear interaction
123 between the wave due to the elastic sheet and the gravity wave near the critical Froude
124 number F_{cr} where their wavelengths approach each other. A steady solution does not exist
125 for a Froude number equal to one of the critical Froude numbers; otherwise, the height of the
126 obstruction should be zero. The new findings are summarized in the Conclusions section.

127 2. Theoretical analysis.

128 A two-dimensional steady flow in a channel with an obstruction on the bottom covered by
129 an elastic sheet representing the ice cover is considered. The obstruction has a characteristic
130 length R , and the thickness of the sheet is \bar{h} . We define a Cartesian coordinate system XY with
131 the origin at the center of the obstruction. The X axis is aligned with the velocity direction
132 of the flow, which has a constant speed U . The Y -axis points vertically upwards. This

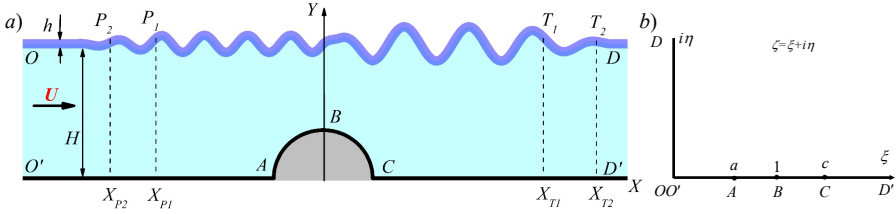


Figure 1: (a) Physical plane and (b) parameter, or ζ -plane.

133 consideration is equivalent to the obstruction moving along the flat bottom of the channel
 134 with velocity U in the opposite direction. A definition sketch of the coordinate system
 135 is shown in Figure 1a. The liquid is inviscid and incompressible, and the flow is assumed to be
 136 irrotational, thus allowing us to use a potential flow model.

137 The obstruction and the bottom downstream are assumed to have an arbitrary shape, which
 138 is defined by the function $Y_b(S)$, where S is the arc length coordinate, or by the slope of the
 139 bottom, $\delta_b = dY/dS$,

$$140 \quad \delta_b(X) = \arctan \frac{dY_b}{dX}.$$

141 We introduce the complex velocity potential, $W(Z) = \Phi(X, Y) + i\Psi(X, Y)$, which consists
 142 of the velocity potential $\Phi(X, Y)$ and the stream function $\Psi(X, Y)$. Here, $Z = X + iY$. The
 143 boundary value problem for the velocity potential can be written as follows:

$$144 \quad \nabla^2 \Phi = 0, \quad \nabla^2 \Psi = 0, \quad (2.1)$$

145 in the liquid domain;

$$146 \quad \frac{\partial \Phi}{\partial Y} = \frac{\partial \Phi}{\partial X} \frac{dY_b}{dX}, \quad \Psi = 0, \quad (2.2)$$

147 on the bottom of the channel $Y_b = Y_b(X)$;

$$148 \quad \rho \frac{V^2}{2} + \rho g Y + p_{ice}(X) + p_{ext}(X) = \rho \frac{U^2}{2} + \rho g H + p_\infty, \quad (2.3)$$

149 which is the dynamic boundary condition at the ice/liquid interface, $Y = Y(X)$. Here,
 150 $V = |\nabla \Phi|$ is the velocity magnitude, $p_{ice}(X)$ is the hydrodynamic pressure at the ice/liquid
 151 interface and $P_\infty = P_a + \rho_i g h$ is its value at infinity; p_a is the atmospheric pressure, ρ_i is the
 152 density of ice, h is the thickness of the ice sheet, and g is the gravity acceleration, $p_{ext}(X)$
 153 is the external pressure applied to the elastic sheet on the intervals $X_{P2} < X < X_{P1}$ and
 154 $X_{T1} < X < X_{T2}$ to provide a waveless interface far upstream and downstream; this will be
 155 discussed in the following.

156 The sought-for solution has the limit $Y(X)_{X \rightarrow -\infty} = H$, where H is the depth of the channel.
 157 The flow is steady; therefore, the value of the stream function at the interface is constant and
 158 equal to the flowrate across the channel

$$159 \quad \Psi = UH; \quad (2.4)$$

160 and the far field condition

$$161 \quad \nabla \Phi \rightarrow U, \quad X \rightarrow -\infty, \quad 0 \leq Y \leq H. \quad (2.5)$$

162 To complete the formulation of the boundary-value problem (2.1) - (2.3), an equation in
 163 the hydrodynamic pressure at the ice/liquid interface is needed. The elastic sheet is modeled

164 using the Cosserat theory of hyperelastic shells (Plotnikov and Toland (2011))

$$165 \quad p_{ice} = D' \left(\frac{d^2\kappa}{dS^2} + \frac{1}{2}\kappa^3 \right) + p_a, \quad (2.6)$$

166 where $D' = \frac{Eh^3}{12(1-\nu^2)}$ is the flexural rigidity of the elastic sheet, κ is the curvature of the
 167 interface, $E = 5.0GPa$ is Young's modulus, and $\nu = 0.3$ is Poisson's ratio. Equation (2.6)
 168 corresponds to the assumption that the elastic sheet is inextensible and is not prestressed. It
 169 should be noted that the difference between the Cosserat theory and the Kirchhoff - Love
 170 plate model, in which the cube of the curvature term in (2.6) is omitted, is quite small due to
 171 a small curvature of the ice sheet before it starts breaking.

172 The interactions between the obstruction, the flow, and the elastic sheet may generate
 173 waves that extend to both upstream and downstream infinity. However, the solutions with
 174 waves extending to upstream infinity are physically meaningless because they do not satisfy
 175 the radiation condition, which requires that there be no energy coming from infinity (Binder,
 176 Vanden-Broeck and Dias (2009)). To satisfy the radiation condition, or make the interface
 177 waveless far upstream, we apply an external pressure on the interval P_1P_2 (see Figure 1a),
 178 which can be located as far as necessary to avoid its effect on the flow near the obstruction,

$$179 \quad p_{ext} = C_d V \frac{dV}{dX}, \quad (2.7)$$

180 where the coefficient C_d characterizes the wave attenuation on the interval P_1P_2 ; it linearly
 181 increases from zero at point P_1 to some value $C_{up} > 0$ at point P_2 and then remains constant.

182 Now we recall that potential flows of an ideal fluid are reversible, i.e., changing the
 183 direction of the inflow velocity has no effect on the results. Alternatively, the flow region can
 184 be mirrored about the y-axis without reversing the velocity direction. Therefore, to make our
 185 solution reversible, it is also necessary to provide a waveless interface far downstream.
 186 Similarly, the external pressure (2.7) is applied on the interval T_1T_2 downstream. The
 187 coefficient C_d changes from zero at point T_1 to some value $C_d = C_{dw}$ at point T_2 and
 188 then remains constant. The same wave attenuation technique was used by Semenov (2021)
 189 for a similar problem, but with an infinite water depth.

190 To solve the problem, it is convenient to nondimensionalize the variables. The velocity U
 191 and the depth of the channel H are used as the reference quantities. Specifically, $x = X/H$
 192 and $y = Y/H$, $s = S/H$, the thickness of the ice sheet h is replaced with $h^* = h/H$, the bottom
 193 profile $y_b(x) = Y_b(X)/H$, and the interface profile $y(x) = Y(x)/H$. The velocity potential Φ
 194 and the stream function Ψ are also normalized to the product UH . The normalized variables
 195 are denoted as $\phi = \Phi/UH$ and $\psi = \Psi/UH$. With these normalizations, the value of the
 196 stream function on the bottom of the channel is $\psi = 0$, and the value of the stream function
 197 at the interface is $\psi = 1$.

198 The nondimensionalized dynamic boundary condition (2.3) takes the form

$$199 \quad v^2 = 1 - \frac{2(y-1)}{F^2} - 2D \left(\frac{d^2\kappa}{dS^2} + \frac{1}{2}\kappa^3 \right) - \frac{2C_d}{H} v \frac{dv}{ds}, \quad (2.8)$$

200 where

$$201 \quad v = |\nabla\phi| = V/U, \quad E_b = \frac{D'}{\rho g H^4}, \quad D = \frac{E_b}{F^2}, \quad \kappa = \frac{d\delta}{ds},$$

202 and

$$203 \quad F = \frac{U}{\sqrt{gH}} \quad (2.9)$$

204 is the Froude number based on the depth of the channel, $\delta = \arcsin(dy/ds) = \beta + \pi$ is the

205 angle between X -axis and the unit tangential vector τ oppositely directed to the velocity
 206 direction β . Equation (2.8) contains the velocity magnitude along the interface v and the wave
 207 elevation y with its derivatives, which will be related in the following through the derived
 208 expression for the complex potential.

209 2.1. Dispersion relation

210 We examine a steady sine-like waving interface of small steepness δ_0 , or the slope of the
 211 interface can be represented as

$$212 \quad \delta(s) = \Re[\delta_0 e^{i k H s}], \quad (2.10)$$

213 where kH is the nondimensional wave number. Upon differentiating equation (2.8) in the arc
 214 length coordinate s , we obtain:

$$215 \quad v^2 \frac{d \ln v}{ds} = - \left(\frac{1}{F^2} + D(kH)^4 \right) \delta. \quad (2.11)$$

216 For the case without an ice sheet ($D = 0$), equation (2.11) becomes

$$217 \quad v^2 \frac{d \ln v}{ds} = - \frac{\delta}{F^2} = - \frac{kH}{\tanh kH} \delta, \quad (2.12)$$

218 where we used the relation between the Froude and wave numbers for free surface gravity
 219 waves in a channel of depth H (Kochin, Kibel and Roze (1964)). We assume that the velocity
 220 along the interface behaves in the same as for the free surface case. From equations (2.11)
 221 and equation (2.12), we obtain the dispersion equation, which coincides, in particular, with
 222 that in the papers Greenhill (1886), Page and Părău (2014)

$$223 \quad \frac{kH}{\tanh kH} = \frac{1}{F^2} + D(kH)^4. \quad (2.13)$$

224 The number of real roots of equation (2.13) depends on the value of the constant D and
 225 the Froude number F . It can have no roots, two roots, or one root. These cases correspond
 226 to a subcritical flow (no roots), $F < F_{cr}$, a channel subcritical and ice supercritical flow (two
 227 roots), $F_{cr} < F < 1$, and a channel supercritical flow ($F > 1$).

228 The wave number versus the Froude number obtained from the solution of equation (2.13)
 229 is shown in Figure 2 for various thicknesses of the ice sheet. It can be seen that without an
 230 ice sheet ($h = 0$) each Froude number $F < 1$ corresponds to one wave number. It tends to
 231 zero as the Froude number $F \rightarrow 1$. In the presence of an ice sheet, there is a minimal, or
 232 critical Froude number F_{cr} , for which the solution of the dispersion equation exists. In the
 233 range $F_{cr} < F < 1$, there are two wave numbers, k_{gr} and k_{ice} corresponding to the gravity
 234 and elastic waves; the wave number $k_{ice} > k_{gr}$, or the elastic wave is shorter than the gravity
 235 wave. This range of the Froude number corresponds to the ice supercritical and channel
 236 subcritical flows. The larger the ice thickness, the smaller the wave number k_{ice} , and the
 237 critical Froude number $F_{cr} \rightarrow 1$. Thus, the interval $F_{cr} < F < 1$, in which both the gravity
 238 and the elastic wave may appear, reduces. For $F > 1$, or for the channel supercritical flows,
 239 there is one root due to the elastic sheet. Since for $F > 1$ the perturbations in the channel
 240 cannot extend upstream, and we may expect the elastic wave extending downstream. Usually,
 241 the dispersion equation 2.13 relates a wave frequency (or phase speed of a monochromatic
 242 wave moving in still water) to the wavenumber: $\omega^2 = k^2 U^2$. In the present case, $U^2 = F^2 g H$;
 243 therefore, the frequency ω and the Froude number are related as $\omega^2 = k^2 F^2 g H$.

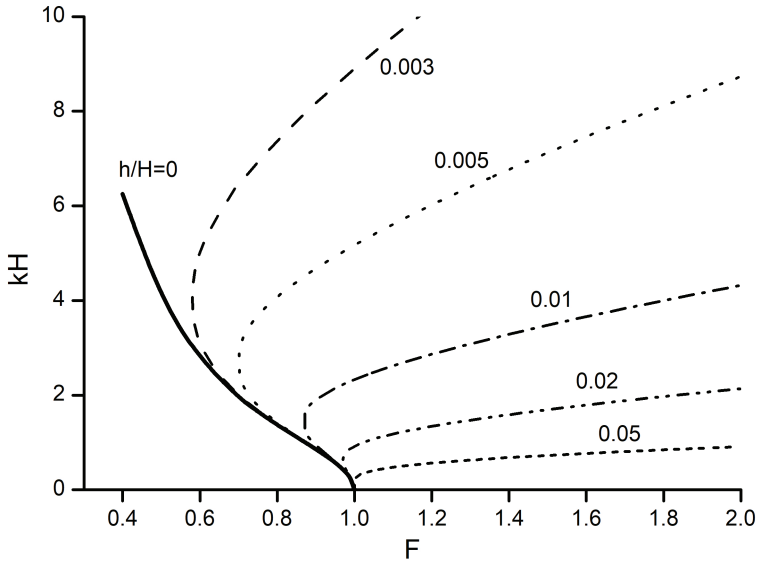


Figure 2: Wave number vs. Froude number for different thicknesses of the ice sheet, h/H .

244

2.2. Integral hodograph method

245 Finding the complex potential of the flow, $w = w(z)$, directly is a complicated problem since
 246 the boundary of the flow region is unknown in advance. Instead, Joukovskii (1890) and
 247 Michell (1890) proposed to introduce an auxiliary parameter plane, or ζ -plane, which was
 248 typically chosen as the upper half-plane. Then, they considered two functions, which were
 249 the complex potential w and the function

$$250 \quad \omega = -\ln\left(\frac{1}{v_0} \frac{dw}{dz}\right) = \ln \frac{v}{v_0} - i\beta, \quad (2.14)$$

251 both functions of the parameter variable ζ . Here, v and β are the velocity magnitude and
 252 direction, respectively; v_0 is the magnitude of the velocity on the free surface, which is
 253 assumed to be constant. When $w = w(\zeta)$ and $\omega(\zeta)$ are derived, the velocity and the flow
 254 region can be obtained in parametric form as follows:

$$255 \quad \frac{dw}{dz} = \exp[-\omega(\zeta)], \quad z(\zeta) = z_0 + \int_0^\zeta \frac{dw}{d\zeta'} / \frac{dw}{dz} d\zeta', \quad (2.15)$$

256 where the function $z(\zeta)$ is called the mapping function.

257 The Joukovskii - Michell method is capable to solve free surface problems for flows over
 258 polygon-shaped bodies and a constant velocity on the free surface/interface (without gravity,
 259 surface tension, etc.). In this case the functions $\omega(\zeta)$ and $w(\zeta)$ form polygon-shaped domains
 260 and can be found applying the Schwarz-Christoffel integral to find their conformal mapping
 261 into the upper half-plane.

262 An additional complexity arises when the slope of the body varies along the body contour
 263 or the velocity magnitude on the free surface/interface varies due to gravity, surface tension,
 264 etc. On these parts of the flow boundary, the boundary conditions are of different types: on
 265 the solid part of the boundary, the velocity direction is determined by the slope of the body;
 266 on the free surface/interface, the velocity magnitude can be obtained from the Bernoulli
 267 equation. This is a so-called mixed boundary-value problem for a complex function.

268 If the upper half-plane is chosen as the region of the parameter variable and the whole real
 269 axis corresponds to the free surface or the body surface, then Schwarz's integral formula or
 270 Cauchy's integral formula can be applied to determine the desired complex function. This
 271 approach was applied by Forbes and Schwartz (1982) for solving free-surface flow over
 272 a semicircular obstruction. In order to use Cauchy's integral formula, they introduced an
 273 image flow symmetric about the x -axis and were able to formulate a uniform boundary-
 274 value problem for the complex function $d\zeta/dw$. By using Cauchy's integral formula and the
 275 dynamic boundary condition, they obtained an integro-differential equation in the complex
 276 function $d\zeta/dw$.

277 In this paper, we use a different integral formula (Semenov and Iafrati (2006), Semenov
 278 and Cummings (2007)) that allows us to determine a complex function based on the values
 279 of its argument and magnitude given on the real and the imaginary axis of the first quadrant,
 280 respectively. Therefore, we chose the first quadrant as the region of the parameter variable
 281 $\zeta = \xi + i\eta$ (instead of a half-plane) shown in Figure 1b. The parameter region corresponds
 282 to the liquid domain in the physical plane $z = x + iy$ shown in Figure 1a: the real axis
 283 corresponds to the bottom of the channel, and the imaginary axis corresponds to the interface.
 284 The conformal mapping theorem allows us to arbitrarily choose the location of three points
 285 $O(O')$ ($\zeta = 0$) B ($\zeta = 1$) and $D(D')$ ($\zeta = \infty$), as shown in 1b. Then, the locations of points
 286 A ($\zeta = a$) and C ($\zeta = c$) are unknown and have to be determined using additional physical
 287 considerations.

288 The complex velocity function on the bottom of the channel and that at the interface are
 289 unknown a priori. At this stage, we assume that these functions are known as functions of
 290 the parameter variables: $v(\eta) = |dw/dz|$ is known as a function of the coordinate η along the
 291 imaginary axis in the ζ -plane; $\chi(\xi) = \arg(dw/dz)$ is a known function of the coordinate ξ
 292 along the real axis of the first quadrant in the ζ -plane. These functions will be determined
 293 later using the dynamic and kinematic boundary conditions at the interface and on the bottom,
 294 respectively. Using the above definitions, we can write the following boundary-value problem
 295 for the complex velocity function:

$$296 \quad \left. \frac{dw}{dz} \right|_{\zeta=i\eta} = v(\eta), \quad 0 \leq \eta < \infty, \quad (2.16)$$

$$297 \quad \arg \left(\left. \frac{dw}{dz} \right|_{\zeta=\xi} \right) = \chi(\xi), \quad 0 \leq \xi < \infty. \quad (2.17)$$

299 By using Chaplygin's singular point method (Gurevich (1965), §5 Chapter 1), the following
 300 integral formula can be obtained for solving the mixed boundary value problem (2.16) and
 301 (2.17) (Semenov and Iafrati (2006)):

$$302 \quad \frac{dw}{dz} = v_\infty \exp \left[\frac{1}{\pi} \int_0^0 \frac{d\chi}{d\xi} \ln \left(\frac{\zeta + \xi}{\zeta - \xi} \right) d\xi - \frac{i}{\pi} \int_0^\infty \frac{d \ln v}{d\eta} \ln \left(\frac{\zeta - i\eta}{\zeta + i\eta} \right) d\eta + i\chi_\infty \right], \quad (2.18)$$

303 where $v_\infty = \lim_{\eta \rightarrow \infty} v(\eta)$ and $\chi_\infty = \lim_{\xi \rightarrow \infty} \chi(\xi)$. An alternative way of derivation of the above
 304 integral formula is presented by Semenov and Cummings (2007). It can easily be verified
 305 that for $\zeta = \xi$ the argument of the function dw/dz is the function $\chi(\xi)$, while for $\zeta = i\eta$ the
 306 magnitude of dw/dz is the function $v(\eta)$, i.e. the boundary conditions (2.16) and (2.17) are
 307 satisfied.

308 The argument of the complex velocity is determined by the slope of the bottom, δ_b , or
 309 $\chi(\xi) = -\delta_b(\xi)$, which at points A and C undergoes a step change due to the corners at points
 310 A and C as can be seen in Figure 1a. We introduce a continuous function $\gamma(\xi)$ that changes

311 from the value $\gamma(a) = 0$ at point A, ($\xi = a$), to the value $\gamma(c) = -\pi$ at point C, ($\xi = c$), and
 312 further may vary continuously along the bottom,

$$313 \quad \chi(\xi) = \begin{cases} 0, & 0 < \xi < a, \\ -\pi/2 - \gamma(\xi), & a \leq \xi \leq c, \\ -\pi - \gamma(\xi), & c < \xi < \infty. \end{cases} \quad (2.19)$$

314 The function $\chi(\xi)$ has two jumps: at point A, $\Delta_A = -\pi/2$ and at point C, $\Delta_C = -\pi/2$. The
 315 function $\gamma(\xi)$ differs from the function $\delta_b(\xi)$ only by a constant; therefore, $d\gamma/d\xi = d\delta_b/d\xi$.
 316 Substituting Eq.(2.19) into Eq.(2.17), evaluating the integrals over the step changes of the
 317 function $\chi(\xi)$, and using $d\gamma/d\xi = d\delta_b/d\xi$, we obtain the expression for the complex velocity
 318 as

$$\frac{dw}{dz} = v_0 \sqrt{\frac{a - \zeta}{a + \zeta} \frac{c - \zeta}{c + \zeta}} \exp \left[-\frac{1}{\pi} \int_a^\infty \frac{d\delta_b}{d\xi} \ln \left(\frac{\xi - \zeta}{\xi + \zeta} \right) d\xi - \frac{i}{\pi} \int_0^\infty \frac{d \ln v}{d\eta} \ln \left(\frac{i\eta - \zeta}{i\eta + \zeta} \right) d\eta \right]. \quad (2.20)$$

319 where $v_0 = 1$ is the velocity magnitude at point O. Here, we used $\arg(\zeta - i\eta) = \arg(i\eta - \zeta) - \pi$
 320 for the second integral.
 321

322 2.3. Derivative of the mapping function, dz/dw

323 On the bottom of the channel the stream function $\psi \equiv 0$, and at the interface $\psi \equiv 1$ as it
 324 follows from the boundary conditions (2.2) and (2.4), while the potential varies from $-\infty$ to
 325 $+\infty$. Thus, the domain of the complex potential $w = \phi + i\psi$ is the infinite strip $-\infty < \phi < \infty$
 326 of unit width, $0 \leq \psi \leq 1$. Due to the simplicity of the domain of w , we can use conformal
 327 mapping to immediately write the complex potential w as a function of the parameter variable
 328 ζ

$$329 \quad w(\zeta) = \frac{2}{\pi} \ln \zeta. \quad (2.21)$$

330 The complex potential (2.21) is a logarithmic function of ζ , or ζ exponentially depends
 331 on the complex potential $w = \phi + i\psi$. The arc length coordinates $s_b \sim \phi$ and $s \sim \phi$ along the
 332 bottom and the interface, respectively. This causes difficulties in computations for a length
 333 of the computational region larger than $5H$. We can resolve the logarithmic singularity if we
 334 eliminate the parameter variables ζ , ξ and η from equation (2.20) using the expressions:

$$335 \quad \left. \begin{aligned} \zeta &= \exp(\pi w/2), & -\infty \leq \phi \leq \infty, & 0 \leq \psi \leq 1, \\ \eta &= \exp(\pi \phi/2), & -\infty \leq \phi \leq \infty, & \psi = 1, \\ \xi &= \exp(\pi \phi/2), & -\infty \leq \phi \leq \infty, & \psi = 0. \end{aligned} \right\} \quad (2.22)$$

336 By substituting (2.22) into (2.20), we obtain the complex velocity as a function of the
 337 complex potential w , the inverse function of which is the derivative of the mapping function,
 338 $z = z(w)$:

$$339 \quad \frac{dz}{dw} = \frac{1}{v_0} \sqrt{\frac{a + e^{w'}}{a - e^{w'}} \frac{c + e^{w'}}{c - e^{w'}}} \exp \left[\frac{1}{\pi} \int_{\phi'_A}^\infty \frac{d\delta_b}{d\phi'} \ln \left(\frac{e^{\phi'} - e^{w'}}{e^{\phi'} + e^{w'}} \right) d\phi' \right. \\ \left. + \frac{i}{\pi} \int_{-\infty}^\infty \frac{d \ln v}{d\phi'} \ln \left(\frac{ie^{\phi'} - e^{w'}}{ie^{\phi'} + e^{w'}} \right) d\phi' \right], \quad (2.23)$$

341 where $w' = \pi w/2$ and $\phi' = \pi\phi/2$. The integrals containing functions

$$342 \quad \ln\left(\frac{e^{\phi'} - e^{w'}}{e^{\phi'} + e^{w'}}\right), \quad \ln\left(\frac{ie^{\phi'} - e^{w'}}{ie^{\phi'} + e^{w'}}\right).$$

343 exponentially decay as the difference $|\phi' - w'|$ increases. The integration of equation (2.23)
344 along $-\infty < \phi < \infty$, $\psi = 1$, in the w -plane gives the interface OD ; its integration along
345 $-\infty < \phi < \infty$, $\psi = 0$ gives the bottom surface. The parameters $a = \exp(\pi\phi_A/2)$ and
346 $c = \exp(\pi\phi_C/2)$. The potentials ϕ_A and ϕ_C , and the functions $\delta_b(\phi)$ and $v(\phi)$ are unknown
347 and have to be determined from physical considerations and the boundary conditions.

348 2.4. Integro-differential equations in the functions $\delta_b(\phi)$

349 By using the derivative of the mapping function (2.23) we can obtain the arc length coordinate
350 s_b as a function of the potential ϕ :

$$351 \quad s_b(\phi) = \int_0^\phi \frac{ds_b}{d\xi} d\phi'. \quad (2.24)$$

352 where

$$353 \quad \frac{ds_b}{d\phi} = \left| \frac{dz}{dw} \right|_{w=\phi} = \frac{1}{v_0} \sqrt{\left| \frac{a + e^{\phi'} c + e^{\phi'}}{a - e^{\phi'} c - e^{\phi'}} \right|} \exp \left\{ \frac{1}{\pi} \int_{\phi_A}^{\infty} \frac{d\delta_b}{d\phi''} \ln \left| \frac{e^{\phi''} - e^{\phi'}}{e^{\phi''} + e^{\phi'}} \right| d\phi'' \right. \\ 354 \quad \left. + \frac{1}{\pi} \int_{-\infty}^{\infty} \frac{d \ln v}{d\phi''} \left[\pi - 2 \tan^{-1} (e^{\phi'' - \phi'}) \right] d\phi'' \right\}, \quad (2.25)$$

355 and $\phi' = \pi\phi/2$.

356 The bottom shape is given by the slope of the bottom, $\delta_b = \delta_b(s_b)$. By making the change
357 of the variables $s_b = s_b(\phi)$ we obtain the following integro-differential equation in the
358 function $\delta_b(\phi)$:

$$359 \quad \frac{d\delta_b}{d\phi} = \frac{d\delta_b}{ds} \frac{ds_b}{d\phi}, \quad (2.26)$$

360 where $ds_b/d\phi$ is determined from the above equation, which also contains the function
361 $d\delta_b/d\phi$. The parameters ϕ_A and ϕ_C are determined from the given arc length of the
362 obstruction ABC . In view of equation (2.24):

$$363 \quad s_{AB} = s_b(\phi_A), \quad s_{BC} = s_b(\phi_C), \quad (2.27)$$

364 where s_{AB} and s_{BC} are the arc lengths of the parts AB and BC of the obstruction.

365 2.5. Determination of the function $v(\phi)$

366 The velocity magnitude at the interface is determined using the dynamic boundary condition
367 (2.8), which contains the interface shape $y(s)$ and curvature with its higher derivatives. The
368 ice/liquid interface is obtained by integrating the derivative of the mapping function (2.23)
369 along the upper side of the strip in the w -plane, or $w = \phi + i$, it takes the form

$$370 \quad x(\phi) + iy(\phi) = x_O + iH + \int_{-\phi^*}^{\phi} \left(\frac{dz}{dw} \right)_{w=\phi+i} d\phi, \quad (2.28)$$

371 where the coordinate of point x_O is obtained by integrating the derivative of the mapping
372 function (2.23) along the lower side of the strip in the w -plane, which corresponds to the

373 bottom of the channel:

$$374 \quad x_O = \int_0^{-\phi^*} \left(\frac{dz}{dw} \right)_{w=\phi} d\phi. \quad (2.29)$$

375 Here, $-\phi^*$ and ϕ^* are the lower and the upper boundary of the computational region; the
376 channel in the physical plane is truncated, and the flow outside the computational region,
377 $|\phi| > \phi^*$, is assumed to be uniform. The arc length coordinate along the interface is

$$378 \quad s(\phi) = \int_0^\phi \frac{d\phi}{v(\phi)}. \quad (2.30)$$

379 It would be possible to determine the slope of the interface using the derivative of the
380 mapping function (2.23),

$$381 \quad \delta(\phi) = \Im \left[\ln \left(\frac{dz}{dw} \right)_{w=\phi+i} \right], \quad (2.31)$$

382 and then evaluate the curvature of the interface and its first and second derivatives by
383 differentiating the equation

$$384 \quad \kappa = \frac{d\delta}{ds} = \frac{d\delta}{d\phi} \frac{d\phi}{ds}. \quad (2.32)$$

385 However, when differentiating the function $\delta(\phi)$ with respect to ϕ , the order of singularity
386 in the integrand of the second integral in equation (2.23) increases. By substituting the
387 y - coordinate of the interface and the second derivative of the curvature into the dynamic
388 boundary condition (2.8), we obtain a very complicated hypersingular integral equation in
389 the function $v(\phi)$, whose numerical solution requires special treatments.

390 Instead of solving the hypersingular integral equation, we use another numerical method
391 based on the spline approximation of the interface to evaluate its curvature and higher
392 derivatives. In discrete form, the solution is sought on two fixed sets of points: a set $-\phi^* <$
393 $\phi_j < \phi^*$, $j = 1, \dots, N$ corresponding to the bottom of the channel and a set $-\phi^* < \phi_i < \phi^*$,
394 $i = 1, \dots, M$ corresponding to the interface; both sets of points ϕ_j and ϕ_i monotonically
395 increase.

396 We chose a fifth-order spline, which provides continuous derivatives along the interface up
397 to the fourth derivative appearing in the pressure coefficient due to the ice sheet

$$398 \quad y(s) = y_k + a_{1,k}(s - s_{k-1}) + \dots + a_{n,k}(s - s_{k-1})^n, \quad s_{k-1} < s < s_k, \quad k = 1, \dots, \bar{K}. \quad (2.33)$$

400 where nodes $s_k = s_{i(k)}$ and $y_k = y_{i(k)}$, $i(k) = 4k - 3$, $k = 1, \dots, \bar{K}$, $\bar{K} = M/4$, are chosen as
401 every 4th point on the set of the discrete points $s_i = s(\phi_i)$ and $y_i = y(\phi_i)$ determined from
402 equations (2.28) and 2.30. The curvature and its derivatives are obtained by differentiating
403 Eq. (2.31):

$$404 \quad \delta = \arcsin y', \quad \kappa = \frac{y''}{\sqrt{1 - y'^2}}, \quad \frac{d\kappa}{ds} = \frac{y' y''^2 - y'''(y'^2 - 1)}{(1 - y'^2)^{3/2}}, \dots$$

405 By applying the dynamic boundary condition (2.8) at the points ϕ_k , $k = 1, \dots, \bar{K}$, we can
406 obtain the following system of nonlinear equations

$$407 \quad G_k(\bar{V}) = c_{pk}(\bar{V}) - c_{pk}^{ice}(\bar{V}) = 0, \quad k = 1, \dots, \bar{K}, \quad (2.34)$$

408 where $\bar{V} = (v_1, \dots, v_{\bar{K}})^T$ is the vector of the unknown velocities v_k at the nodes s_k ;

$$409 \quad c_{pk}(\bar{V}) = 1 - v_k^2 - \frac{2[y_k(\bar{V}) - 1]}{F^2} - C_{dvk} \left(\frac{dv}{dx} \right)_k, \quad (2.35)$$

$$410 \quad c_{pk}^{ice}(\bar{V}) = 2D \left[\left(\frac{d^2\kappa}{ds^2} \right)_k + \frac{1}{2}\kappa_k^3 \right]. \quad (2.36)$$

411 are the hydrodynamic pressure coefficient and the pressure coefficient due to the elastic
412 sheet, respectively. The wave **attenuation** intervals are chosen to be $x_{P1} - x_{P2} = 2\lambda_{gr}$ and
413 $x_{T2} - x_{T1} = 3\lambda_{gr}$, where λ_{gr} is the wavelength determined from the dispersion relation. The
414 coefficients C_{up} and C_{dw} are chosen in the interval from $0.2\lambda_{gr}$ to $0.4\lambda_{gr}$ to effectively damp
415 both the elastic and the gravity wave upstream and downstream, respectively.

416 The system of equations (2.34) is solved using Newton's method. The Jacobian of the
417 system is evaluated numerically using the central difference with $\Delta v_k = 10^{-8}$. At each
418 evaluation of the function $G_k(\bar{V})$, the integro-differential equation (2.26) together with
419 equations (2.25) and (2.27) is solved using the method of successive approximations, **which**
420 **in discrete form becomes**

$$421 \quad \frac{(\Delta\delta_b)_j^{(m+1)}}{\Delta\phi_j} = \frac{\delta_b(s_{bj}^{(m)}) - \delta_b(s_{b(j-1)}^{(m)})}{\Delta\phi_j}, \quad (2.37)$$

422 where the arc length along the body, $s_b^{(m)} = s_b^{(m)}(\phi_j)$ is evaluated using (2.24) with
423 $(\Delta\delta_b)_j^{(m)} / \Delta\xi_j$ known at iteration m . The iteration process converges very fast. After 5
424 to 10 iterations, the error is below a prescribed tolerance of 10^{-6} . The parameters a and c
425 are obtained as

$$426 \quad a = \exp(\pi/2\phi_A) \quad c = \exp(\pi/2\phi_C), \quad (2.38)$$

427 where ϕ_A and ϕ_C are determined from equation (2.27). From 5 to 20 iterations are necessary
428 to get a converged solution. All solutions, say \bar{V}^* , reported here satisfied the condition

$$429 \quad \frac{1}{\bar{K}} \sum_1^{\bar{K}} |G_k(\bar{V}^*)| < 10^{-6}, \quad (2.39)$$

430 which is considered as giving a sufficiently accurate solution of the nonlinear equations.

431 At the first iteration, the functions $v(\phi)$, $s_b(\phi)$, and $\delta_b(\phi)$ and the parameters ϕ_A and ϕ_C
432 are specified as follows: $v^{(1)}(\phi) \equiv 1$, $s_b^{(1)}(\phi) = \phi$, $\phi_A^{(1)} = s_{AB}$, $\phi_C^{(1)} = s_{BC}$, and

$$433 \quad \delta_b^{(1)}(\phi) = \begin{cases} \pi/2, & -\infty < \phi \leq \phi_A, \\ \pi/2 - \pi(\phi - \phi_A)/(\phi_C - \phi_A), & \phi_A \leq \phi \leq \phi_C, \\ -\pi/2, & \phi_C \leq \phi < \infty. \end{cases} \quad (2.40)$$

434 Then, the next iteration starts with solving the integro-differential equation (2.26).

435 3. Results and discussion

436

3.1. Numerical approach

437 The number of nodes on the bottom and at the interface is chosen in the ranges $200 < N < 400$
438 and $400 < M < 4000$, respectively, based on the requirement to provide at least 12 nodes
439 s_k within the shorter wavelength and to get a reasonably accurate converged solution. **On**

440 a Precision Tower desktop T7920, the computational time varies from a few minutes for
441 $M = 400$ to about 30min for $M = 4000$.

442 The integrals appearing in Eq. (2.23) are evaluated analytically using points of discretiza-
443 tion of the real and the imaginary axis of the first quadrant in the ζ -plane, $\xi_j = \exp(\pi\phi_j/2)$
444 and $\eta_i = \exp(\pi\phi_i/2)$, and a linear interpolation of the functions $\delta_b(\xi)$ on the intervals
445 (ξ_{j-1}, ξ_j) , and the function $\ln v(\eta)$ on the intervals (η_{i-1}, η_i) :

$$446 \quad \frac{1}{\pi} \int_{\xi_{j-1}}^{\xi_j} \frac{d\delta_b}{d\xi} \ln \left(\frac{\xi - \zeta}{\xi + \zeta} \right) d\xi = \Delta\delta_{bj} a_j(\zeta) \quad (3.1)$$

$$448 \quad \frac{1}{\pi} \int_{-\eta_{i-1}}^{\eta_i} \frac{d \ln v}{d\eta} \ln \left(\frac{i\eta - \zeta}{i\eta + \zeta} \right) d\eta = \Delta(\ln v)_i b_i(\zeta) \quad (3.2)$$

449 where

$$450 \quad \Delta\delta_{bj} = \delta_b(\xi_j) - \delta_b(\xi_{j-1}),$$

$$451 \quad a_j(\zeta) = \frac{1}{\pi\Delta\xi_j} \int_{\xi_{j-1}}^{\xi_j} \ln \left(\frac{\xi - \zeta}{\xi + \zeta} \right) d\xi,$$

$$452 \quad \Delta(\ln v)_i = \ln v_i - \ln v_{i-1} = \ln \frac{v_i}{v_{i-1}},$$

$$453 \quad b_i(\zeta) = \frac{1}{\pi\Delta\eta_i} \int_{\eta_{i-1}}^{\eta_i} \ln \left(\frac{i\eta - \zeta}{i\eta + \zeta} \right) d\eta.$$

454 The integral in the above equation can be easily evaluated, and the result is a nonsingular
455 expression for the functions $a_j(\zeta)$ and $b_j(\zeta)$. By substituting (2.22) into (3.1) and (3.2) we
456 can evaluate the integrals in equation (2.23).

457 3.2. Verification of the numerical approach

458 For verification purposes we compare the results predicted by the present nonlinear solution
459 with the nonlinear theory Page and Părău (2014) based on the boundary integral method for
460 the case of a hydraulic fall. They considered the hydraulic fall solution for which the depth
461 of liquid upstream is greater than downstream. The flow is assumed to be uniform in the far
462 field as $|x| \rightarrow \infty$, with a constant depth H and velocity U downstream and a constant depth
463 $H_{up} > H$ and velocity $U_{up} < U$ upstream of the obstruction on the bottom of the channel.

464 Applying Bernoulli's equation in the far fields $|x| \rightarrow \pm\infty$ and using the conservation mass
465 equation, the parameters upstream and downstream are related in nondimensional form as
466 follows (Dias and Vanden-Broeck (2004)):

$$467 \quad \frac{1}{2} - \frac{1}{2}\gamma^{*2} + \frac{1}{F^2} - \frac{1}{F^2\gamma^{*2}} = 0, \quad (3.3)$$

468 where $\gamma^* = U_{up}/U$

469 Page and Părău (2014) considered a cosine-squared profile of the bottom of the channel
470 including two obstructions as follows

$$471 \quad y_b(x) = \begin{cases} 2A_1 \cos^2 \left(\frac{\pi(x+x_1)}{2L_1} \right), & -L_1 < x < x+x_1 < L_1, \\ 2A_2 \cos^2 \left(\frac{\pi x}{2L_2} \right), & -L_2 < x < x < L_2, \\ 0, & \text{otherwise.} \end{cases} \quad (3.4)$$

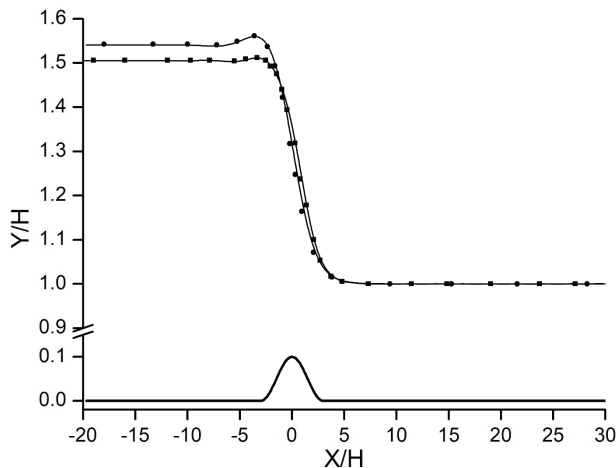


Figure 3: Hydraulic fall profiles over a single submerged obstruction of: height $2A_2 = 0.1$ and length $L_2 = 6$; $E_b = 0.5$, $F = 1.367$, $\gamma^* = 0.649$ (solid circles); $E_b = 0.1$, $F = 1.345$, $\gamma^* = 0.664$ (solid squares); lines and symbols correspond to the present solution and Page and Părau (2014), respectively.

472 The heights and half-lengths of the submerged obstructions are defined by $2A_i$ and L_i
 473 ($i = 1, 2$), respectively. The separation constant x_1 describes the central position of the
 474 additional obstruction. In the case of just a single submerged obstruction, A_1 is taken to be
 475 zero.

476 The hydraulic fall profiles over an obstruction predicted by the present solution are
 477 compared with the results by Page and Părau (2014) in Figure 3 for two cases with Froude
 478 number $F = 1.367$ and $F = 1.345$. It can be seen that the results predicted by the present
 479 method and that by Page and Părau (2014) coincide.

480 An additional verification is performed for the case of two obstructions on the bottom.
 481 In the absence of a thin ice sheet, placing an additional obstruction downstream of the
 482 hydraulic fall in the pure gravity case can result in a train of trapped waves between the two
 483 obstructions (Dias and Vanden-Broeck (2004)). Page and Părau (2014) predicted trapped
 484 waves in the presence of an ice sheet. The interface profile and the bottom shape for the case
 485 of an additional obstruction centred at $x = 20$ is shown in Figure 4 for the present solution
 486 (solid line) and Page and Părau (2014) (dashed line and solid symbols). The Froude number
 487 F_{cr} is found as part of the solution using the additional condition of the absence of a wave
 488 downstream. If this condition is not applied and Froude number $F > 1$ is given, a wave
 489 downstream of the second obstruction can be observed. This will be discussed later in the
 490 following. An agreement between the present results and those by Page and Părau (2014)
 491 verifies the calculation code.

492

3.3. Subcritical flows, $F < F_{cr}$

493 For Froude numbers $F < F_{cr}$, equation (2.13) has only complex roots, which correspond to
 494 decaying perturbations of the interface caused by the obstruction on the bottom. In Figure
 495 5a, the interface profiles for obstruction height $R/H = 0.2$ and thickness of the ice elastic
 496 sheet $h/H = 0.01$ are shown for different Froude numbers approaching the critical Froude
 497 $F_{cr} = 0.864$. It can be seen that the interface shape is symmetric about the y -axis and the
 498 wave decays; the trough of the wave is located just above the obstruction, and it gets deeper
 499 as the Froude number approaches the critical value. This situation is different from that for

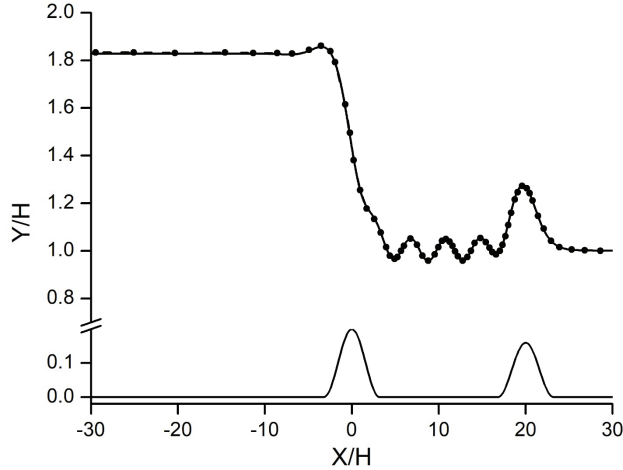


Figure 4: Trapped wave at the ice/liquid interface (the solid line corresponds to the present calculations; the dashed line with symbols corresponds to Page and Părau (2014)) for the bottom profile including two obstructions (blue line): $2A_2 = 0.2$ and width $2L_2 = 6.4$ and an additional obstacle with $2A_1 = 0.16$ and $2L_1 = 6.4$ placed at $x_1 = 20$; the Froude number $F = 1.5373$ and $\gamma^* = 0.545$ are found as part of the solution, and $E_b = 0.5$.

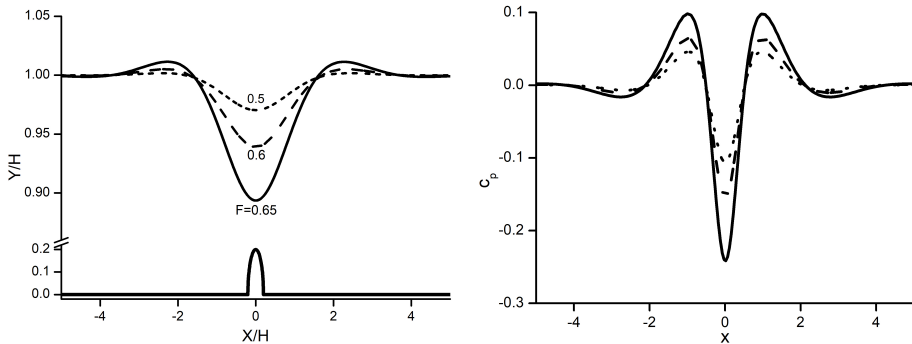


Figure 5: (a) the interface shape and (b) the pressure coefficient along the interface for obstruction height $R/H = 0.2$, ice thickness $h/H = 0.01$, and a subcritical flow: Froude number $F = 0.65$ (solid line), $F = 0.6$ (dashed line); $F = 0.5$ (dotted line).

500 the free-surface flows without an elastic sheet, for which the free surface is flat upstream
 501 and exhibits a wave downstream of the obstruction. Thus, for $F < F_{cr}$, the elastic sheet
 502 suppresses the waves downstream and perturbs the flow near the obstruction. It was found
 503 that for $R/H = 0.2$ and Froude $0.65 < F < F_{cr}$ the solution fails to converge, or $F = 0.65$ is
 504 the maximal value.

505 The interface profiles for different heights of the obstruction and the maximal value of the
 506 Froude number for each height are shown in Figure 6: for height $R/H = 0.32$ the maximal
 507 Froude number is $F = 0.5$; for $R/H = 0.17$, $F = 0.7$; and for $R/H = 0.06$, $F = 0.83$. As
 508 the height of the obstruction further decreases, the maximum Froude number approaches the
 509 critical Froude number $F_{cr} = 0.864$. It can be seen from Figure 6a that the smaller the height
 510 of the obstruction, the smaller the deflection of the interface corresponding to the onset of
 511 convergence of the solution. Therefore, we can conclude that a large height of the obstruction

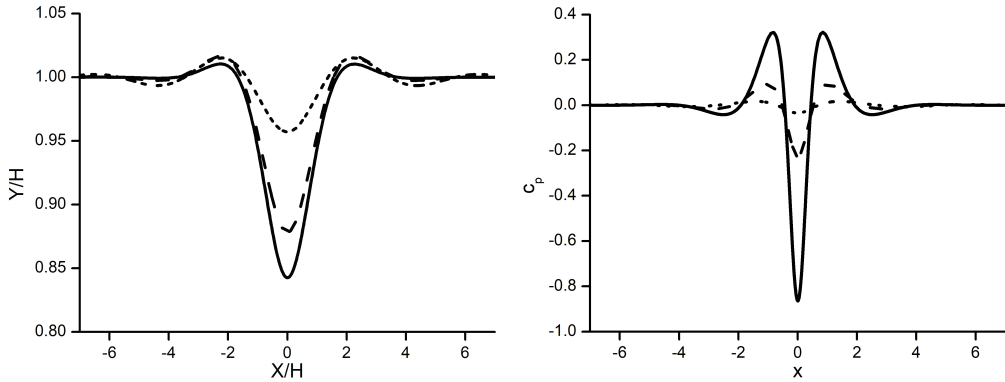


Figure 6: (a) the interface shape and (b) the pressure coefficient along the interface corresponding to the onset of convergence of the subcritical solution for ice thickness $h/H = 0.01$: $F = 0.5$, $R/H = 0.32$ (solid line); $F = 0.7$, $R/H = 0.17$ (solid line); $F = 0.83$, $R/H = 0.06$ (solid line); the critical Froude number $F_{cr} = 0.8636$.

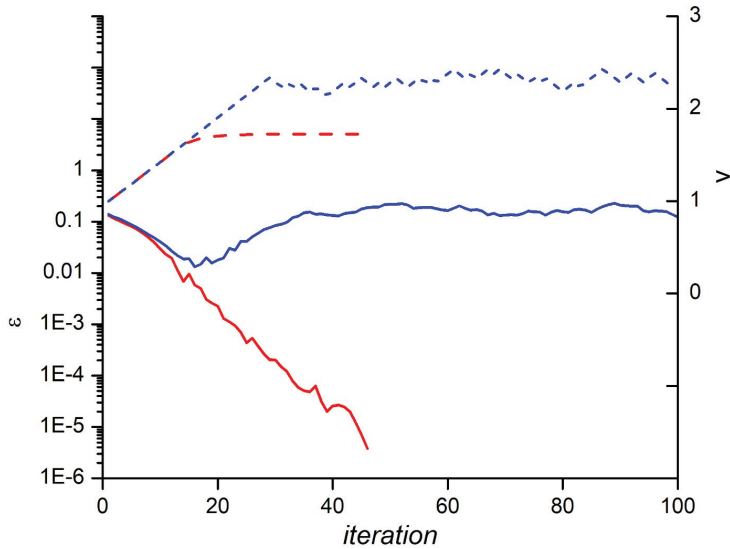


Figure 7: Convergence of the iterations for Froude number $F = 0.5$ and two heights of the obstruction, $R/H = 0.32$ (red line) and $R/H = 0.33$ (blue line); the left axis corresponds to the average error in the dynamic boundary condition (solid lines); the right axis corresponds to the velocity magnitude at the trough (dashed lines).

512 or a large deflection of the ice/liquid interface themselves do not prevent the convergence of
513 the solution.

514 The behavior of the average error (2.39) and the velocity magnitude at the trough is shown
515 in Figure 7 (the left and the right axis, respectively) for two slightly different heights of the
516 obstructions. The initial velocity at the interface is set to $v(\phi) \equiv 1$. At the beginning of
517 the iterations, the velocity magnitude at the trough increases linearly for both cases due to
518 the given restriction of the velocity increment. For $R/H = 0.32$ the average error gradually
519 decreases and the velocity at the trough tends to some value, while for $R/H = 0.33$ both the
520 error and velocity magnitude oscillate without any tendency to converge.

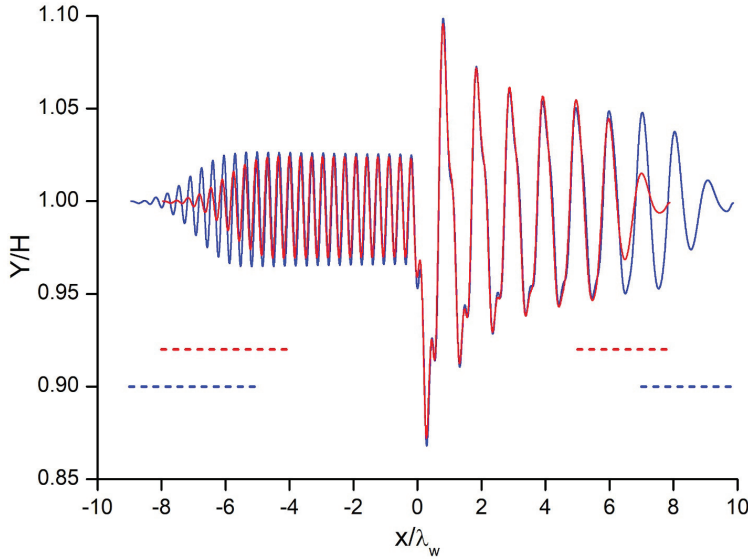


Figure 8: The ice/water interface for Froude number $F = 0.8$, obstruction height $R/H = 0.11$, and different lengths of the computational region: $16\lambda_{gw}$ (red line); $19\lambda_{gw}$ (blue line). The dashed lines indicate the length and location of the **attenuation** zones.

521

3.4. Ice supercritical – channel subcritical flows, $F_{cr} < F < 1$

522

In this range of the Froude number, the dispersion relation (2.13) has two real roots: one root, k_{gw} , is the wave number corresponding to the gravity wave (since its value is close to the wave number corresponding to the free surface gravity waves), and the other root, k_{ice} , is caused by the elastic wave. Both waves may extend to infinity downstream and upstream. In order to examine how the introduced **attenuation** regions affect the solution, we compare in Figure 8 the ice/liquid interfaces corresponding to two cases: for the first case, the computational region starts at $x_{P2} = -8\lambda_{gw}$ and ends at $x_{T2} = 8\lambda_{gw}$ with **attenuation** zone length $L_{P1P2} = 4\lambda_{gw}$ and $L_{T1T2} = 3\lambda_{gw}$; for the second case, it starts at $x_{P2} = -10\lambda_{gw}$ and ends at $x_{T2} = 9\lambda_{gw}$ with the same length of the **attenuation** zones. For both cases, the **attenuation** coefficients are $C_{up} = 0.14\lambda_{gw}$ and $C_{dw} = 0.4\lambda_{gw}$; for the both cases, the Froude number $F = 0.8$, $R/H = 0.11$ and the ice thickness $h/H = 0.005$, for which the critical Froude number $F_{cr} = 0.6935$. From the dispersion relation (2.13), the wave numbers are obtained: $k_{gw} = 1.4252$ ($\lambda_{gw} = 4.409$) and $k_{ice} = 4.0757$ ($\lambda_{ice} = 1.542$). The number of nodes of the spline \bar{K} is chosen to provide at least 12 nodes within the shorter ice wave λ_{ice} . Then, the total number of nodes for the first case, $x_{T2} - x_{P2} = 16\lambda_{gw}$, is obtained as $\bar{K} = 12 * 16\lambda_{gw}/\lambda_{ice} \approx 550$, and the total number of discretization points at the interface is $M = 4\bar{K} = 2200$. For the second case, the length of the computational region $x_{T2} - x_{P2} = 19\lambda_{gw}$, and $\bar{K} \approx 650$ and $M = 2600$.

540

The red and blue solid lines in Figure 8 correspond to the first case and the second case, respectively. The dashed lines indicate the location and the length of the **attenuation** zones for each case. It can be seen that the red and blue lines coincide in the range of $-4 < x\lambda_{gw} < 5$ where the **attenuation** term in the dynamic boundary condition (2.8) is equal to zero ($C_d = 0$).

544

As the Froude number $F \rightarrow 1$, the ratio $\lambda_{gw}/\lambda_{ice} = k_{ice}/k_{gw} \rightarrow \infty$ since the gravity wave number $k_{gw} \rightarrow 0$. In this case the required number of discretization points also tends to infinity, thus causing computational difficulties. The computational analysis starts with

546

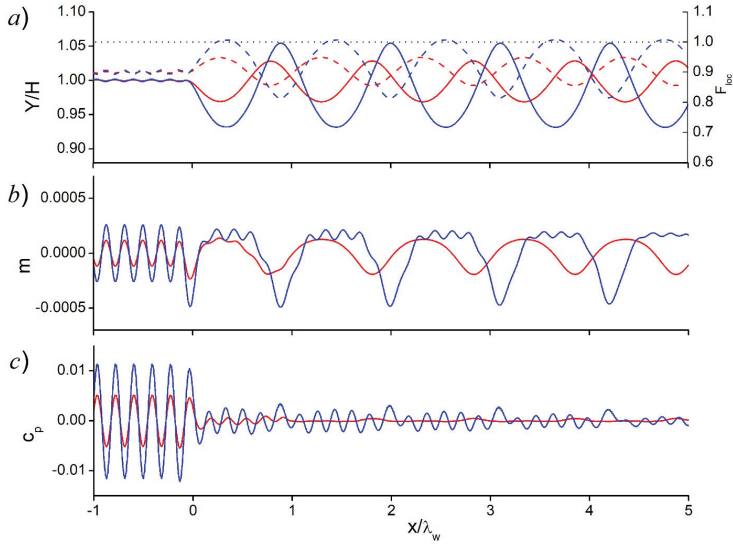


Figure 9: The ice/water interface (a), the bending moment (b), and the pressure coefficient (c) along the interface for Froude number $F = 0.9$, ice thickness $h/H = 0.005$, and obstruction height $R/H = 0.05$ (red line) and 0.07 (blue line), which is the maximal height for which the steady solution is obtained; the critical Froude number $F_{cr} = 0.6935$.

547 $F = 0.9$ and then gradually approaches the critical Froude number $F_{cr} = 0.6935$; the ice
 548 thickness $h/H = 0.005$. The wavelengths of the gravity and the elastic wave are $\lambda_{gw} = 7.250$
 549 and $\lambda_{ice} = 1.344$, and their ratio $\lambda_{gw}/\lambda_{ice} = 5.394$.

550 Figure 9 shows (a) the interface profile, (b) the bending moment, and (c) the pressure
 551 coefficient along the interface for two heights of the obstruction: $R/H = 0.05$ (red line) and
 552 0.07 (blue line), Froude number $F = 0.9$, or $F/F_{cr} = 1.30$, and ice thickness $h/H = 0.005$.
 553 This ratio is relatively large in terms of interaction between the gravity and the elastic wave,
 554 which is quite weak. For $R/H = 0.05$ the interface is almost flat upstream, or the oscillations
 555 of the elastic wave are invisible, although its contribution to the bending moment and the
 556 pressure coefficient is significant. For a larger height, $R/H = 0.07$, the wave amplitude of
 557 the interface upstream becomes visible, but it is still much lower than the amplitude of the
 558 interface downstream corresponding to the gravity wave.

559
$$F_{loc} = \frac{v(x)}{\sqrt{y(x)}} F \quad (3.5)$$

560 The dashed lines correspond to the local Froude number (right axis). It can be seen in
 561 Figure 9a that the local Froude number for $R/H = 0.05$ (red dashed line) does not reach the
 562 channel critical value ($F = 1$), and the period of the wave is close to λ_{gw} predicted by the
 563 dispersion relation. For $R/H = 0.07$, the amplitude of the both the elastic and the gravity
 564 wave increases, and the local Froude number reaches the critical value at the wave trough on
 565 some shot intervals. This means that the flow becomes transcritical on some shot intervals,
 566 thus affecting the wavelength, which is slightly increased. We found that the convergence
 567 of the solution for obstruction height $R/H > 0.07$ is very challenging: the amplitude and
 568 period of the gravity wave further increase, which results in the lowering of the interface and

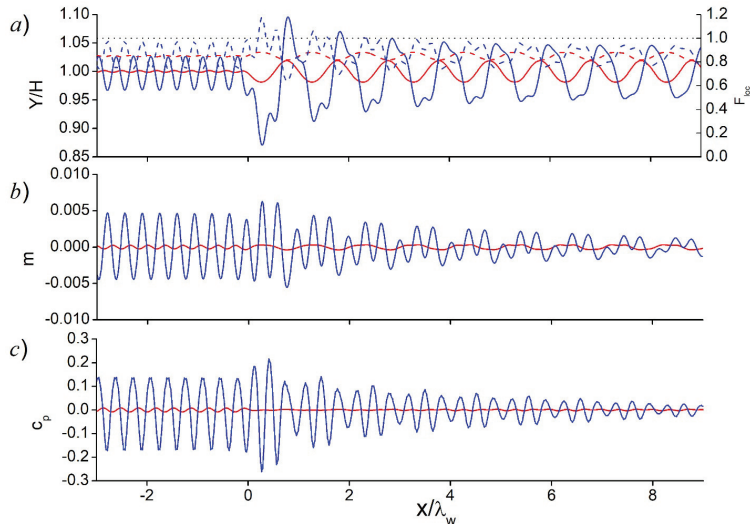


Figure 10: The same as in Figure 9 for $F = 0.8$ and $R/H = 0.05$ (red line) and 0.1 (blue line).

569 a further increase in the velocity at the trough. The supercritical part of the flow becomes
570 larger.

571 The bending moment along the interface is shown in Figure 9b. It can be seen that the
572 amplitudes of the bending moment for the elastic wave upstream and the gravity wave
573 downstream are about the same. For obstruction height $R/H = 0.05$, the bending moment
574 varies quite smoothly both upstream and downstream of the obstruction. For $R/H = 0.07$, the
575 bending moment exhibits sine-like behavior upstream of the obstruction, but downstream we
576 can observe a sharp trough corresponding to the crest at the interface and a flat interval
577 of bending with a small contribution of the elastic wave, which gradually decays. A
578 superposition of the gravity and the elastic wave is clearly seen because the wavelength
579 ratio $\lambda_{gw}/\lambda_{ice} = 5.394$ is relatively large. For a smaller ratio, the interaction of the waves will
580 cause more complicated behaviour of the interface, the bending moment, and the pressure
581 coefficient.

582 The behavior of the pressure coefficient along the interface is shown in Figure 9c. It can be
583 seen that the amplitude of the oscillations upstream is much higher than those downstream.
584 The oscillations of the pressure coefficient due to gravity downstream are so small that they
585 are almost invisible. That is why we can observe downstream only a small contribution
586 caused by the elastic wave, which gradually decays.

587 The results for Froude number $F = 0.8$, or $F/F_{cr} = 1.15$, and two obstruction heights
588 $R/H = 0.05$ and 0.11 are shown in Figure 10. The wavelengths are: $\lambda_{gw} = 4.409$ and
589 $\lambda_{ice} = 1.542$; the ratio $\lambda_{gw}/\lambda_{ice} = 2.86$. For height $R/H = 0.05$, the amplitude of the
590 elastic wave upstream is quite small in comparison with the amplitude of the gravity wave
591 downstream. Both waves exhibit sine-like behavior. For height $R/H = 0.11$, the amplitude
592 of the elastic wave increases, so that the local Froude number approaches the critical value
593 $F = 1$ at the troughs. This causes difficulties in the convergence of the solution for larger
594 heights of the obstruction. The interface downstream exhibits a superposition of the gravity
595 wave and the elastic wave, although the contribution of the latter decays. However, since the
596 wavelengths approach each other, their interaction exhibits more complicated behavior than

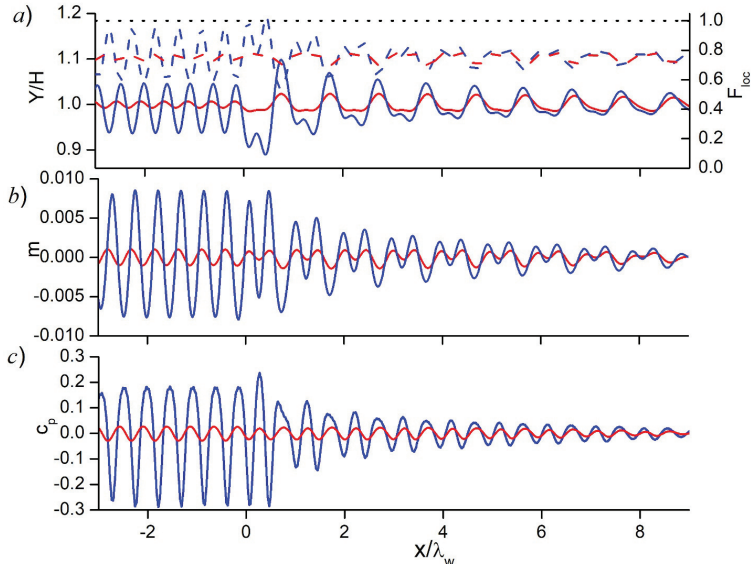


Figure 11: The same as in Figure 9 for $F = 0.75$ and $R/H = 0.05$ (red line) and 0.11 (blue line).

597 that in Figure 9. The bending moment and the pressure coefficient along the interface are
 598 shown in Figures 10b and 10c. For the smaller height of the obstruction, the oscillations
 599 caused by the elastic sheet and gravity can be seen upstream and downstream separately. For
 600 the larger height, the gravity does not affect the oscillations of bending moment upstream,
 601 while the elastic sheet contributes to a superposition of the oscillations downstream to a
 602 larger extent, and its contribution decays downstream slower than in Figure 9b.

603 The results for Froude number $F = 0.75$, or $F/F_{cr} = 1.08$, and two obstruction heights
 604 $R/H = 0.05$ and 0.11 are shown in Figure 11. The wavelengths are: $\lambda_{gw} = 3.537$ and
 605 $\lambda_{ice} = 1.705$; the ratio $\lambda_{gw}/\lambda_{ice} = 2.07$. This case is closer to the critical Froude number
 606 $F_{cr} = 0.6935$, and we can observe a larger amplitude of the interface upstream (due to the
 607 elastic sheet), while the amplitude of the wave downstream (due to gravity) becomes smaller.
 608 Moreover, for $R/H = 0.11$ the amplitude of the elastic wave upstream becomes larger than
 609 the amplitude of the gravity wave far downstream, where the contribution of the elastic wave
 610 decays. The length of the computational region in Figure 11 may not be large enough to see
 611 the interface without any contribution of the elastic wave. When the Froude number further
 612 approaches the critical Froude number F_{cr} , the interaction of the elastic and the gravity wave
 613 gets stronger. This results in a smaller height of the obstruction for which the solution can be
 614 obtained.

615

3.5. Channel supercritical flows, $F > 1$

616 It is well known for free-surface channel flows (Dias and Vanden-Broeck (1989)) that for
 617 the supercritical regime ($F > 1$) there may exist two solutions, one with a smaller height
 618 of the wave crest called the 'perturbed' wave and the other with a higher wave crest called
 619 the soliton wave. The 'perturbed' wave is a solution that is a member of a family of steady
 620 solutions that bifurcate from the uniform stream as the height of the obstruction increases
 621 from zero. The 'soliton' wave is a member of a family of steady solutions that bifurcate from
 622 a solitary wave as the height of the obstruction increases from zero. The families merge at

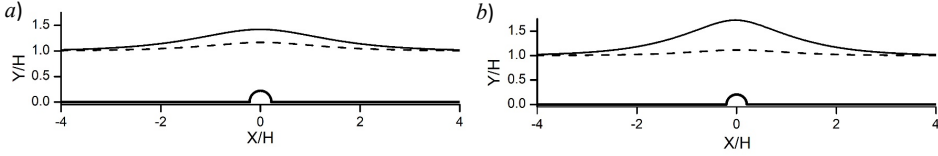


Figure 12: The perturbed wave (dashed line) and the soliton wave (solid line) for a free-surface channel supercritical flow with Froude number (a) $F = 1.2$ and (b) $F = 1.3$; the height of the obstruction $R/H = 0.2$.

623 a fold for some Froude number, F_{min} , which is the minimum Froude number. If a solitary
 624 wave does not exist, then a 'soliton' type solution does not exist either. There is no solution
 625 of any type in the range $1 < F < F_{min}$ (Dias and Vanden-Broeck (1989)).

626 The present method allows one to compute both these cases. The perturbed wave, $y(x)$,
 627 is computed in parametric form using equations (2.28) and (2.29). In order to compute the
 628 soliton wave, we rearrange the obtained free-surface/interface $y(x)$ in such a way as to fit its
 629 maximum value with the given coordinate y_{ms} of the soliton crest,

$$630 \quad \bar{y}(s) = H + C_{ms}(y(s) - H), \quad (3.6)$$

631 where

$$632 \quad C_{ms} = \frac{y_{ms} - H}{y_m - H}, \quad y_m = \max_{s(-\phi^*) < s < s(\phi^*)} [y(s)].$$

633 The unknown coordinate of the soliton crest, y_{ms} , is obtained by solving the equation

$$634 \quad C_{ms}(y_{ms}) = 1, \quad (3.7)$$

635 then the functions $\bar{y}(s)$ and $y(s)$ are coincided.

636 The perturbed (dashed line) and the soliton (solid line) wave for Froude numbers (a)
 637 $F = 1.3$ and (b) $F = 1.2$ are shown in Figure 12; the height of the obstruction $R/H = 0.2$.
 638 For this height, the soliton wave was found in the range of the Froude number $1.2 < F < 1.4$.
 639 As the Froude number approaches the upper limit, the free surface of the soliton wave forms
 640 an angle of 120 degrees at the wave crest (Vanden-Broeck (1987)).

641 In contrast to a channel flow with a free surface or a liquid surface covered by broken ice
 642 Ni et al. (2023), the attempts to find a soliton wave in the presence of an elastic sheet were
 643 unsuccessful. In the following, the analysis of perturbed-type supercritical flows is presented.

644 Figure 13 shows the interface profiles (a), the bending moment (b), and the pressure
 645 coefficient (c) along the interface for the perturbed type of channel supercritical flow.
 646 Froude number $F = 1.2$ is the minimal value for which a converged solution is obtained for
 647 obstruction height $R/H = 0.2$. For this case (red line), the interface reaches its maximum
 648 above the obstruction, and the local Froude number (red dashed line) drops below 1, or the
 649 local flow becomes subcritical. A subcritical flow at some part of the interface may generate
 650 local waves there, which may hinder the convergence of the iterative process. In Figure 13a,
 651 it can be seen that there are no waves upstream (because the flow is supercritical), but there
 652 are small (almost invisible) waves downstream. These waves manifest themselves clearly in
 653 the behavior of the bending moment and the pressure coefficient shown in Figure 13b and
 654 c. The wave amplitudes of the interface, the bending moment, and the pressure coefficient
 655 decrease as the Froude number increases.

656 The ice/water interfaces for thickness $h/H = 0.01$ are shown in Figure 14 for different
 657 Froude numbers. In comparison with the results for $h/H = 0.005$ in Figure 13, the oscillation
 658 of the ice/liquid interface about the perturbed free surface is clearly seen. The wave attenuation

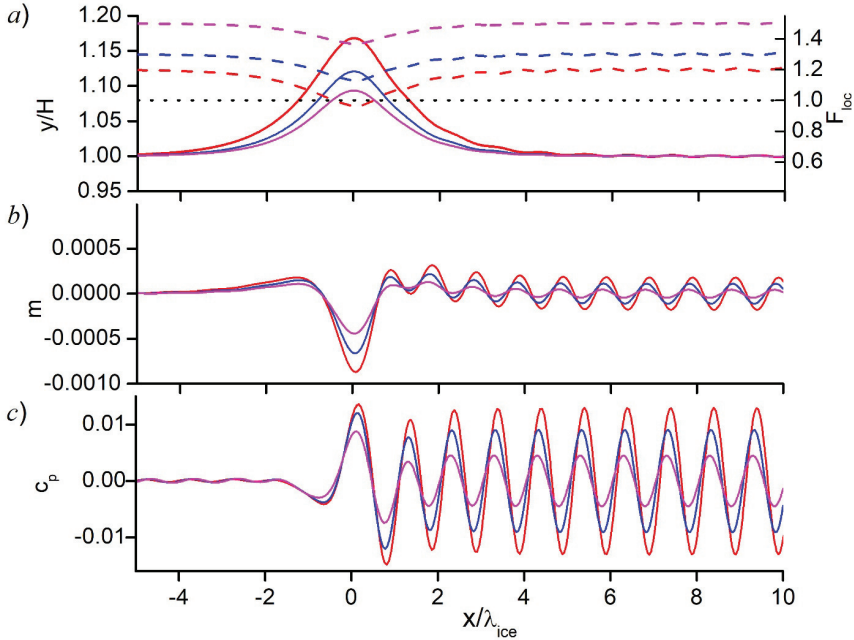


Figure 13: Supercritical flows for Froude numbers $F = 1.2$, $\lambda_{\text{ice}} = 1.0434$, (red), $F = 1.3$, $\lambda_{\text{ice}} = 0.981$, (blue), and $F = 1.5$, $\lambda_{\text{ice}} = 0.882$ (magenta) and $R/H = 0.2$: (a) the ice/water interfaces (solid lines) and the local Froude number (dashed lines); (b) bending moment, and (c) the pressure coefficient.

659 term in the dynamic boundary condition (2.8) was applied far downstream. The sign of the
 660 coefficient C_{dw} was taken negative to provide wave attenuation for the case of the channel
 661 supercritical flows, $F > 1$.

662 The elastic wave for $F_{\text{cr}} < F < 1$ in Figures 9 to 11 propagates upstream, while for $F > 1$
 663 in Figures 13 and 14 it propagates downstream. The wave number k_{ice} in both cases coincides
 664 with that predicted by the dispersion equation (2.13), and it is continuous near $F = 1$ (see
 665 Figure 2). Therefore, one would expect that even at $F > 1$ the elastic wave remains upstream
 666 rather than appearing downstream. Such a case is possible from a mathematical point of view
 667 if we recall that the potential flows of an ideal fluid are reversible, i.e., changing the direction
 668 of the inflow velocity has no effect on the results. Then the elastic wave will propagate
 669 upstream and the downstream flow will be waveless. This is because the flow direction does
 670 not appear in the formulation of the boundary value problem (2.1) - (2.5). The choice of the
 671 ‘correct’ flow direction depends on how the solution corresponds to real observations. Dias
 672 and Vanden-Broeck (2002) studied a generalised hydraulic fall with a free surface. They
 673 found that the radiation condition is satisfied only for waves propagating downstream.

674 Finally, let us justify the results shown in Figure 13 and reffig14, for which the velocity is
 675 directed from left to right. For $F > 1$, a perturbation in the liquid cannot move upstream, so
 676 there is no perturbation of the ice sheet from the liquid, and consequently no wave upstream
 677 is excited.

678 The scaled strain, $\varepsilon_{xx}/\varepsilon_Y$, where $\varepsilon_{xx} = -\frac{1}{2}h\kappa$ is the strain in the floating elastic plate
 679 and ε_Y is the yield strain for the ice estimated as $8 \cdot 10^{-5}$, see Brocklehurst et al. (2011),
 680 is shown in Figure 14 by red lines. For the obstacle with $R/H = 0.2$, the scaled strain is

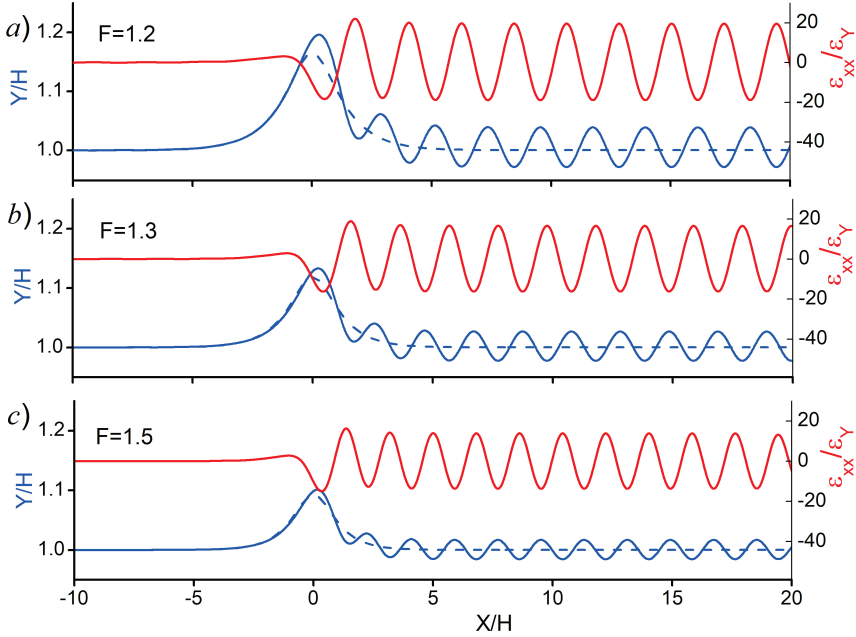


Figure 14: The perturbed supercritical ice/water interfaces (blue solid lines), the free surface without an ice sheet (blue dashed lines) and the scaled strain, e_{xx}/e_Y (red lines) for Froude numbers (a) $F = 1.2$, (b) $F = 1.3$, and (c) $F = 1.5$; ice thickness $h/H = 0.01$, and obstruction height $R/H = 0.2$.

681 less than one in magnitude only well above the obstacle. Formally speaking, the obtained
 682 solution predicts that the continuous ice sheet should be broken starting from $X/H = -2$.
 683 Note that yield strain ε_Y is not used for calculations of the ice elevation. If the ice is less
 684 brittle, which is ε_Y is greater than our estimate, then the ice could be not damaged even for
 685 the conditions of Figure 14. It is understood that the strains in the ice cover are smaller for
 686 smaller obstacles. For given characteristics of the ice cover and a given speed of the current,
 687 we can find the maximal height of the obstacle before the scaled strain $\varepsilon_{xx}/\varepsilon_Y$ exceeds one.
 688 Different characteristics of elastic plate placed on the water above an obstacle, as those used
 689 in the laboratory experiments by Pogorelova et al. (2019) provide different conditions of the
 690 plate damage.

691 The scaled strains in Figure 14 can be well approximated by sinusoidal functions
 692 downstream from the obstacle,

$$693 \quad \frac{e_{xx}}{e_Y} = A \sin \left[kH \frac{x}{H} + \delta \right] + A_0,$$

694 where

$$695 \quad A = 19.2, \quad kH = 2.866, \quad \delta = 2.42, \quad A_0 = 0.30, \quad \text{for } F = 1.2;$$

$$696 \quad A = 16.4, \quad kH = 3.087, \quad \delta = 2.85, \quad A_0 = 0.20, \quad \text{for } F = 1.3;$$

$$697 \quad A = 13.8, \quad kH = 3.485, \quad \delta = 2.94, \quad A_0 = 0.15, \quad \text{for } F = 1.5.$$

698 The non-zero values A_0 indicate that the waves downstream the obstacle are non-linear.
 699 However, the dimensionless wave numbers kH obtained from the numerical solution satisfy
 700 the dispersion equation (2.13) for linear waves with relative accuracy less than 0.4%. The

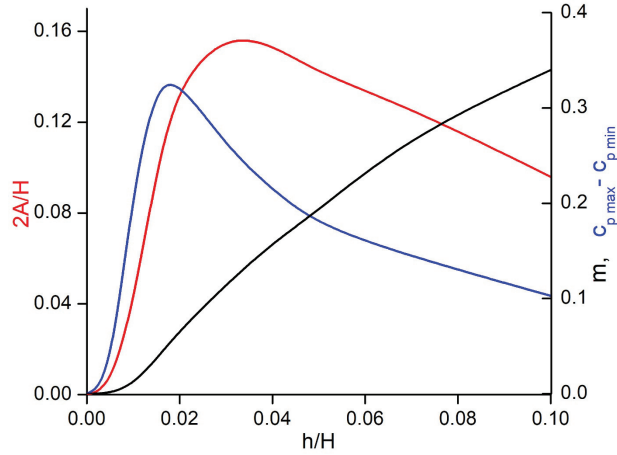


Figure 15: Wave amplitudes of the interface downstream of the obstruction versus the ice sheet thickness h/H : the interface (red, left axis), its bending moment (magenta, right axis), and the pressure coefficient (blue, right axis) for Froude number $F = 1.4$ and obstruction height $R/H = 0.2$.

701 relative difference was calculated as the difference between the left hand side and right hand
 702 side of (2.13) divided by the left hand side and multiplied by 100%.

703 The wave amplitude of the ice/liquid interface, the bending moment, and the pressure
 704 coefficient versus the thickness of the ice sheet are shown in Figure 15 for Froude number
 705 $F = 1.4$ and obstruction height $R/H = 0.2$. In the case of the free surface ($h = 0$), waves
 706 are absent, or the amplitude is equal to zero. For the case of a very large thickness of the ice
 707 sheet, it behaves like a rigid plate; therefore, waves are absent too. Therefore, there exists a
 708 thickness of the ice sheet for which the wave amplitude reaches its maximal value. It can be
 709 seen in Figure 15 that the amplitude of the interface reaches its maximal value at thickness
 710 $h/H = 0.033$, while the pressure coefficient takes its maximal value at $h/H = 0.18$. The
 711 bending moment gradually increases in the range $h/H < 0.1$ presented in the figure. For a
 712 larger ice thickness, computations become challenging because the waves become very long
 713 (see Figure 2) and require too many discretization points, for example, for $h/H = 0.1$ and
 714 $F = 1.4$ $\lambda_{ice}/H = 16.2$.

715 Throughout the analysis of the results discussed in this section starting with the subcritical
 716 flows and ending with the channel supercritical flows, it was shown that the obstruction
 717 height plays an important role: it determines the level of flow nonlinearity and affects the
 718 existence of the solution. It was found that as the Froude number approaches one of the
 719 critical Froude numbers F_{cr} or $F = 1$, the obstruction height corresponding to the onset of
 720 existence of the solution becomes smaller. This is shown in Figure 16 in the Froude number
 721 versus obstruction height plane for two thicknesses of the ice sheet, $h/H = 0.005$ and 0.010 .
 722 The reasons restricting the height of the obstruction near the F_{cr} and $F = 1$ are different: near
 723 the critical value F_{cr} , but $F_{cr} < F$ the lengths of the elastic and the gravity wave approach each
 724 other, and they exhibit a complicated interaction; near the channel critical Froude number,
 725 $F = 1$, the flow downstream becomes transcritical, or it becomes subcritical at the wave crest,
 726 while the flow is supercritical upstream; alternatively, it becomes supercritical at the trough,
 727 while the flow is channel subcritical ($F_1 < 1$) upstream. The larger the height of the obstruction,
 728 the more the Froude number deviates from the critical values. It can also be seen in 16 that

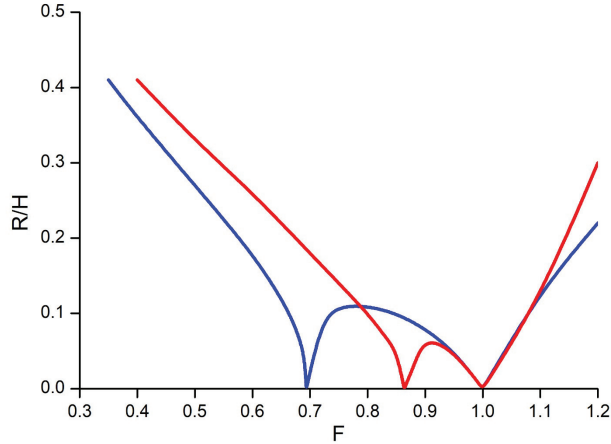


Figure 16: The onset of existence of the steady solution (exists below the line) in the Froude number versus obstruction height plane; the ice thickness $h/H = 0.005$ (blue line) and $h/H = 0.010$ (red line).

729 the region between the two critical Froude numbers, $1 - F_{cr}$, and the maximal height of the
 730 obstruction get smaller.

731 4. Conclusions

732 Fully nonlinear solutions of the flexural-gravity waves in a channel covered by an elastic
 733 sheet are obtained. A case study is presented for a channel of constant depth with a semi-
 734 circular obstruction on the bottom. The integral hodograph method is adopted to solve the
 735 boundary value problem in two steps. At the first step, an expression for the complex velocity
 736 is obtained using the integral formula that solves the mixed boundary value problem for the
 737 first quadrant, which is the chosen parameter region. At the second step, the parameter variable
 738 of the first quadrant is eliminated by using the relation between it and the complex potential
 739 w . Then, the complex potential w is used as the independent variable in the expression for
 740 the derivative of the mapping function, which facilitates the computations in the channel at
 741 larger distances from the obstruction in both directions. A system of integral equations in
 742 the slope of the bottom and the velocity magnitude at the interface is obtained using the
 743 kinematic and dynamic boundary conditions. In discrete form, the problem is reduced to a
 744 system of nonlinear equations in the unknown magnitude of the velocity at the interface,
 745 which is solved numerically using a collocation method. The numerical model is verified by
 746 computing hydraulic fall solutions and comparing the results with those by Page and Părău
 747 (2014).

748 According to the dispersion relation, there are three intervals of the Froude number for
 749 which the interface behaves differently. The first corresponds to the subcritical flows $F < F_{cr}$,
 750 for which the disturbance of the ice/liquid interface caused by the submerged body decays
 751 both in the upstream and in the downstream direction; the second is the ice supercritical
 752 and channel subcritical interval, $F_{cr} < F < 1$, which is characterized by the elastic wave
 753 extending to infinity upstream and the gravity wave extending to infinity downstream; the
 754 third interval corresponds to the channel supercritical flows, $F > 1$, for which the obstruction
 755 generates a hydroelastic wave downstream oscillating about the perturbed free surface wave.

756 It is found that for each Froude number there exists a restriction on the obstruction height for
757 which a converged solution can be obtained.

758 The most complicated behavior of the interface was found for the second range of the
759 Froude number where the two waves caused by the elastic sheet and gravity interact with
760 each other. The gravity wave is observed only downstream, while the elastic wave extends
761 to infinity upstream and some distance downstream of the obstruction. The contribution of
762 the elastic wave to the resulting interface shape decays downstream at a rate that depends on
763 the ratio $\lambda_{gw}/\lambda_{ice}$, or F/F_{cr} . For a relatively large ratio of the wavelengths, the elastic wave
764 decays very fast, and its contribution to the resulting interface can be observed considering
765 only the behavior of the bending moment and the pressure coefficient. As the ratio $\lambda_{gw}/\lambda_{ice}$
766 approaches one, or $F/F_{cr} \rightarrow 1$, the elastic wave weakly decays downstream. The length
767 and amplitude of the waves are about the same; therefore, they exhibit a strongly nonlinear
768 interaction. In order to get a converged solution, the height of the obstruction should be taken
769 small enough.

770 For the channel supercritical flows, $F > 1$, we found a wave caused by the elastic sheet
771 whose wave number agrees with that predicted by the dispersion relation. The wave oscillates
772 about the perturbed free surface solution for the case without an elastic sheet. The amplitude
773 of the wave depends on the thickness of the elastic sheet. It is obvious that there is no wave
774 downstream for the cases $h/H = 0$ (the free surface) and $h/H \rightarrow \infty$ (the rigid plate). From
775 the computations, we found the maximal amplitude of the hydroelastic wave downstream
776 and the thickness of the elastic sheet, $h/H \approx 0.033$, to which it corresponds; the pressure
777 coefficient reaches its maximal value for sheet thickness $h/H \approx 0.018$.

778 The Forbes and Schwartz (1982) found for the free surface flows that there is no solution
779 for Froude number $F = 1$. The present solution confirmed this result for the cases of an
780 elastic sheet and revealed that no solution exists for the critical Froude number F_{cr} .

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