

1 **Fast oxidation of sulfur dioxide by hydrogen**  
2 **peroxide in deliquesced aerosol particles**

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10 **Classification: PHYSICAL SCIENCES: Earth, Atmospheric, and Planetary**

11 **Sciences**

12 **Abstract**

13 Atmospheric sulfate aerosols have important impacts on air quality, climate, and human  
14 and ecosystem health. However, current air-quality models generally underestimate the  
15 rate of conversion of sulfur dioxide (SO<sub>2</sub>) to sulfate during severe haze pollution events,  
16 indicating that our understanding of sulfate formation chemistry is incomplete. This  
17 may arise because the air-quality models rely upon kinetics studies of SO<sub>2</sub> oxidation  
18 conducted in dilute aqueous solutions, and not at the high solute strengths of  
19 atmospheric aerosol particles. Here, we utilize an aerosol flow reactor to perform direct  
20 investigation on the kinetics of aqueous oxidation of dissolved SO<sub>2</sub> by hydrogen  
21 peroxide (H<sub>2</sub>O<sub>2</sub>) using pH-buffered, submicrometer, deliquesced aerosol particles at  
22 relative humidity of 73 to 90%. We find that the high solute strength of the aerosol  
23 particles significantly enhances the sulfate formation rate for the H<sub>2</sub>O<sub>2</sub> oxidation  
24 pathway compared to the dilute solution. By taking these effects into account, our  
25 results indicate that the oxidation of SO<sub>2</sub> by H<sub>2</sub>O<sub>2</sub> in the liquid water present in  
26 atmospheric aerosol particles can contribute to the missing sulfate source during severe  
27 haze episodes.

28 **Keywords: air pollution; Chinese haze; sulfate aerosol; sulfur dioxide oxidation;**  
29 **aerosol kinetics; multiphase chemistry**

30 **Significance**

31 Atmospheric sulfate aerosol particles contribute significantly to poor air quality and  
32 direct forcing of the Earth's climate. However, air pollution and climate models  
33 simulate the formation of sulfate using acid rain chemistry known to be appropriate  
34 only for cloud water conditions. By measuring the oxidation of sulfur dioxide (SO<sub>2</sub>) by  
35 hydrogen peroxide (H<sub>2</sub>O<sub>2</sub>) directly in hygroscopic, pH-buffered aerosol particles with  
36 high solute strength characteristic of many tropospheric conditions, we show that  
37 sulfate formation occurs significantly faster than under the cloudwater conditions  
38 previously explored. In part, ionic strength and general acid catalysis effects drive the  
39 fast chemistry. These results indicate that the H<sub>2</sub>O<sub>2</sub>-driven oxidation of SO<sub>2</sub> in aqueous  
40 aerosol particles can contribute to the missing sulfate source during severe haze  
41 pollution events.

## 42 **Introduction**

43 Sulfate aerosol is an important component of fine particulate matter that impacts air  
44 quality, climate, and human and ecosystem health (1-3). Atmospheric models currently  
45 generate aerosol sulfate either via condensation of H<sub>2</sub>SO<sub>4</sub>, which is formed via gas-  
46 phase oxidation of SO<sub>2</sub>, or via a suite of oxidation processes involving SO<sub>2</sub> dissolved  
47 in cloudwater. Although the aerosol liquid water content (AWC) is generally much  
48 lower than cloud liquid water, it is possible that such aerosol multiphase oxidation  
49 processes may be important in polluted and high relative humidity conditions. However,  
50 it is uncertain whether the kinetics of aqueous oxidation of dissolved SO<sub>2</sub> by different  
51 oxidants investigated in bulk solution with low ionic strength (< 5 molal) are applicable  
52 to the high solute concentration situations that prevail for aerosol particles. Measuring  
53 the kinetics of aqueous oxidation of dissolved SO<sub>2</sub> in aerosol particles is thus critical to  
54 the accurate modeling of aerosol sulfate in the atmosphere.

55       Assessing the rate of aerosol sulfate formation in polluted conditions can evaluate  
56 the atmospheric importance of multiphase oxidation processes. In particular, rapid  
57 sulfate production has been observed during cloud-free, severe haze events in China,  
58 with the peak sulfate mass concentration reaching as high as ~ 300 μg m<sup>-3</sup> (4-8).  
59 However, current air-quality models that include gas-phase oxidation of SO<sub>2</sub> by the  
60 hydroxyl radical (9) and aqueous oxidation of dissolved SO<sub>2</sub> by hydrogen peroxide  
61 (H<sub>2</sub>O<sub>2</sub>) (10), O<sub>3</sub> (10), O<sub>2</sub> catalyzed by transition metal ions (TMI, i.e., Fe (III) and Mn  
62 (II)) (11-14), methyl hydrogen peroxide (15), and peroxyacetic acid (15) cannot capture  
63 these high levels of aerosol sulfate (5, 16, 17), indicating that our understanding of

64 sulfate formation chemistry is fundamentally incomplete. Oxidation of dissolved SO<sub>2</sub>  
65 by NO<sub>2</sub> may be important if the aerosol pH is high (4, 6) and inclusion of a hypothetical  
66 heterogeneous oxidation process in aerosol particles can greatly improve the model  
67 performance (17). Overall, the formation mechanism of the missing sulfate source  
68 remains unclear and controversial (4, 6, 18-26).

69 Oxidation experiments with high solute strength aerosol face significant  
70 challenges due to the need for online measurement of the reaction kinetics using aerosol  
71 particles, and the necessity for good control of AWC and aerosol pH (27). Sulfate  
72 formation rates for many aqueous SO<sub>2</sub> oxidation pathways involving O<sub>3</sub>, O<sub>2</sub>+TMI, and  
73 NO<sub>2</sub> are strongly pH-dependent (6) and are subject to 1 to 2 orders of magnitude change  
74 if the pH changes by 1 unit because the solubility and effective Henry's law constant of  
75 SO<sub>2</sub> positively depend on pH (28). This sensitivity of sulfate formation rates to pH  
76 poses experimental challenges in controlling aerosol pH because product hydrogen ions  
77 (H<sup>+</sup>) will perturb the aerosol pH. As an exception, the rate of aqueous oxidation of  
78 dissolved SO<sub>2</sub> by H<sub>2</sub>O<sub>2</sub> is largely pH-independent for pH above 2 because the effects  
79 arising from the pH dependence of the SO<sub>2</sub> solubility and the reaction rate constant  
80 offset each other (10). This characteristic makes the SO<sub>2</sub>-H<sub>2</sub>O<sub>2</sub> reaction a useful system  
81 to isolate the effects of solute strength from aerosol pH on the sulfate formation rate.

82 Here, we study pH-buffered submicrometer, deliquesced aerosol particles in an  
83 aerosol flow tube to create high solute strengths, enabling direct investigation of the  
84 kinetics of aqueous oxidation of dissolved SO<sub>2</sub> by H<sub>2</sub>O<sub>2</sub> in aerosol particles. Six types  
85 of seed aerosols were investigated, with aerosol pH buffered at 2.3 to 4.8 as calculated

86 using the E-AIM model (29, 30) and a Pitzer activity coefficient model (31): A) a  
87 mixture of sodium chloride (NaCl)/malonic acid/sodium bimalonate (molar ratios  
88 20:1:1, 6:1:1, 2:1:1 and aerosol pH 2.3–2.5), B) NaCl/sodium bimalonate/sodium  
89 malonate (molar ratios 20:1:1, 6:1:1, 2:1:1 and aerosol pH 4.8), C) sodium nitrate  
90 (NaNO<sub>3</sub>)/malonic acid/sodium bimalonate (molar ratio 20:1:1 and aerosol pH 2.8), D)  
91 NaNO<sub>3</sub>/sodium bimalonate/sodium malonate (molar ratio 20:1:1 and aerosol pH 4.0),  
92 E) malonic acid/sodium bimalonate (molar ratio of 1:1 and aerosol pH 2.8), and F)  
93 sodium bimalonate/sodium malonate (molar ratio of 1:1 and aerosol pH 3.9).

94 In part, the individual aerosol particles components were chosen to be  
95 representative of species found in the atmosphere. More importantly, they satisfy the  
96 demands of the experiment (see detailed explanation in *SI Appendix*, section S1) by  
97 providing different aerosol pH and different AWC (32), and by enabling explicit  
98 examination of the effects of aerosol pH and solute strength on the sulfate formation  
99 rate. Aerosol mass spectrometry (33) (AMS) quantitatively characterized the  
100 composition of seed aerosols and the sulfate that forms. A scanning mobility particle  
101 sizer (SMPS) was used to measure particle-size distributions and to determine the AWC.  
102 All experiments were performed at 21 to 25 °C and high relative humidity (RH) (73 to  
103 90%) to ensure that the seed aerosol particles are deliquesced; see Methods and *SI*  
104 *Appendix*, section S1 for details on the experimental conditions, choice of aerosol  
105 systems to study, instrument operation, and data analysis.

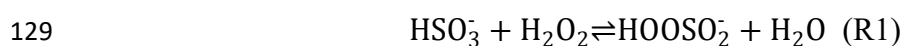
106 The overall goal of this work is to measure the sulfate formation rates on aerosol  
107 particles with high solute concentrations, to enable comparison with the literature

108 parameters that have previously been obtained in bulk solutions. The results will  
109 provide an improved quantitative understanding of sulfate aerosol formation during  
110 severe haze pollution events.

## 111 **Results**

112 The kinetics of aqueous oxidation of dissolved SO<sub>2</sub> by H<sub>2</sub>O<sub>2</sub> in a mixture of NaCl and  
113 malonic acid buffer aerosol particles (molar ratio of 20:1:1) at pH 2.5 are shown in Fig.  
114 1. Results for the other aerosol types are qualitatively similar and shown in *SI Appendix*  
115 Fig. S3. The partial pressures of SO<sub>2</sub> and H<sub>2</sub>O<sub>2</sub> were kept constant and in excess in each  
116 experiment so that the aerosol sulfate that forms is solely dependent on the reaction  
117 time. The dissolved sulfate concentrations, [SO<sub>4</sub><sup>2-</sup>] (molality units), show strong linear  
118 correlations ( $r^2 > 0.96$ ) with the reaction time (Figs. 1A and C). As well, the slopes of  
119 the sulfate formation rate versus initial SO<sub>2</sub> and H<sub>2</sub>O<sub>2</sub> concentrations using log-log plots  
120 (Figs. 1B and D) are close to unity ( $1.03 \pm 0.14$  and  $1.19 \pm 0.03$ , respectively),  
121 suggesting first-order reactions in dissolved SO<sub>2</sub> and H<sub>2</sub>O<sub>2</sub>. Except for some data points  
122 associated with the particles containing NaCl or NaNO<sub>3</sub>, the estimated buffer capacity  
123 is higher than the amount of H<sup>+</sup> formed (see *SI Appendix*, section S1). It is possible that  
124 HCl or HNO<sub>3</sub> evaporate from the particles under those conditions, removing acidity.  
125 Nevertheless, the strong correlations in Fig. 1A and C show no signs of a slower  
126 reaction when a large amount of sulfate forms.

127 Overall, these observations are consistent with the assumed mechanism for the  
128 reaction (34):





132 As well, it is known that weak acids, HX, can act as proton donors to promote the  
 133 reaction through general acid catalysis:



135 The overall rate expression (34, 35) valid for  $\text{pH} > 2$  is:

136 
$$\frac{d\text{SO}_4^{2-}}{dt} = (k + k_{\text{HX}} [\text{HX}][\text{H}^+]^{-1}) K_{\text{a}1} H_{\text{SO}_2} P_{\text{SO}_2} H_{\text{H}_2\text{O}_2} P_{\text{H}_2\text{O}_2}$$
 (1)

137 where  $k = k_{\text{R}3} K_{\text{R}1} K_{\text{R}2}$ , and  $k_{\text{R}3}$  is a reaction rate coefficient,  $K_{\text{R}1}$  and  $K_{\text{R}2}$  are  
 138 thermodynamic equilibrium constants,  $K_{\text{a}1}$  is the thermodynamic dissociation constant  
 139 of  $\text{H}_2\text{SO}_3$ , and  $H_i$  and  $P_i$  represent the Henry's law constant and partial pressure of  
 140 species  $i$ , respectively.  $k_{\text{HX}}$  is the overall rate constant for the general acid catalysis  
 141 mechanism. We note that Eq. 1 is only applicable to ideal solutions. Known  
 142 modifications to Eq. 1 for ionic strength effects are presented in *SI Appendix*, Table S4.

143 To convert the measured sulfate formation rate to units of  $\mu\text{g m}^{-3} \cdot \text{h}^{-1}$ , which are  
 144 used for comparison to field measurements, the following equation was used:

145 
$$\frac{d\text{SO}_4^{2-}}{dt} (\mu\text{g m}^{-3} \text{h}^{-1}) = 3600 \text{ s h}^{-1} \cdot 96 \text{ g mol}^{-1} \cdot \frac{d\text{SO}_4^{2-}}{dt} (\text{molal s}^{-1}) \cdot \frac{\text{AWC}}{\rho_w}$$
 (2)

146  
 147 where AWC is in units of  $\text{mg m}^{-3}$  and  $\rho_w$  is the water density in  $\text{kg L}^{-1}$ . Note that in  
 148 previous studies (6, 19), an equivalent expression was employed with the sulfate  
 149 formation rate in units of  $\text{M s}^{-1}$ . While this is appropriate for dilute conditions, for  
 150 solutions with high ionic strength the aerosol liquid water volume is substantially  
 151 smaller than the total aerosol volume and so Eq. 2 is more accurate. As a result, we



152 determine the  $\text{SO}_2$ - and  $\text{H}_2\text{O}_2$ -dependent sulfate formation rates in molal  $\text{s}^{-1}$ , as  
153 illustrated by the slopes in Fig. 1A and C.

154 Fig. 2A and B show the ratio of the measured sulfate formation rates to the  
155 modeled ones for all six types of seed aerosol as a function of total solute concentration  
156 and ionic strength, respectively. Modeled sulfate formation rates were calculated based  
157 on the literature parameters of  $\text{SO}_2$  oxidation by  $\text{H}_2\text{O}_2$  obtained in bulk solutions,  
158 without taking the effects of ionic strength and general acid catalysis into account; see  
159 *SI Appendix*, section S2 for details. Overall, the measured sulfate formation rates are  
160 higher than the modeled values. For the mixed NaCl and malonic acid buffer aerosol  
161 particles with low ionic strengths ( $I < 7$  molal) and low concentrations of malonic acid  
162 ( $< 0.6$  molal) (Exp 1 through 7 and 12), the measured sulfate formation rates are  $2.1 \pm$   
163  $0.1 - 2.9 \pm 0.6$  times larger than the modeled ones. Overall, we conclude for these  
164 particles types that the pH does not affect the kinetics between pH values of 2.3 and 4.8.  
165 This is consistent with the bulk solution kinetics in the literature for solution pH values  
166 above 2.

167 For the more concentrated solutions ( $I > 7$  molal), the ratio of the measured to the  
168 modeled sulfate formation rate displays strongly increasing trends with increasing  
169 malonic acid concentration and ionic strength as shown in Fig. 2B. The ratio is as high  
170 as  $33 \pm 4$  to  $51 \pm 9$  for the malonic acid buffer aerosol particles at pH 3.9 with ionic  
171 strength of around 14 molal. We believe there are at least 2 effects driving the increase:  
172 general acid catalysis by malonic acid and ionic strength effects. In addition, there is

173 possibility of salting-in effects of SO<sub>2</sub> and H<sub>2</sub>O<sub>2</sub> at high solute strengths, including the  
174 impact of high buffer concentrations, that are difficult to estimate.

175 To illustrate the general acid effects, the measured sulfate formation rate increases  
176 with malonic acid concentration for the mixed NaCl and malonic acid buffer aerosol  
177 particles at pH 2.3 to 2.4 at relatively constant ionic strengths of 3.9 and 6.6 molal (*SI*  
178 *Appendix*, Fig. S6 and Table S1, Exp# 13 through 18).

179 To isolate the ionic strength effect, it is necessary to account for the general acid  
180 catalyzed enhancement. To do this, we linearly fit the two  $k_{HX}$  reaction rate coefficients  
181 that we measured (*SI Appendix*, Fig. S6) as a function of ionic strength (*SI Appendix*,  
182 Fig. S7A), and use this fit to account for the general acid effect according to Eq. 1. The  
183 dependence of  $k_{HX}$  on ionic strength arises because  $k_{HX}$  is dependent on the  $pK_a^*$  of the  
184 general acid (*SI Appendix*, Fig. S7B) and the first  $pK_a^*$  of malonic acid depends on  
185 ionic strength (*SI Appendix*, Fig. S7C), where  $K_a^*$  is the stoichiometric dissociation  
186 constant. From this analysis we conclude that the general acid catalysis effect on the  
187 systems with highest ionic strength (14 molal) is negligible and that the significant  
188 increase of the sulfate formation rate is largely due to the effect of ionic strength on the  
189 proton-catalyzed mechanism and, potentially, the salting-in of reactants as well. This is  
190 consistent with previous studies at lower ionic strengths (~5 molal NaCl) that have  
191 found enhancements in the proton-catalysis reaction rate coefficient  $k$  (34) and the  
192 Henry's law constant of H<sub>2</sub>O<sub>2</sub> (36) (*SI Appendix*, Fig. S5). Referring to the reaction  
193 mechanism above, the strong dependence of the rate constant on ionic strength arises  
194 by lowering the stoichiometric dissociation constant of HOOSO<sub>2</sub>H (reverse of Reaction

195 R2) (34). In particular, at high ionic strength there will be less water to stabilize the ions  
196 into which HOOSO<sub>2</sub>H dissociates.

197 Relative to modeled rates in pure water, we plot the dependence of the  
198 enhancement factor of only the proton-catalyzed sulfate formation rate on ionic strength  
199 in Fig. 2C, i.e., the general acid catalyzed sulfate formation rate was subtracted from  
200 the measured sulfate formation rate to yield the proton-catalyzed sulfate formation rate  
201 (see details in *SI Appendix*, section S3). This factor reflects the overall effects of ionic  
202 strength on the reaction rate coefficient  $k$ , Henry's law constants of H<sub>2</sub>O<sub>2</sub> and SO<sub>2</sub>, and  
203 the first stoichiometric dissociation constant of H<sub>2</sub>SO<sub>3</sub>. We note that some data points  
204 are negative after correcting for the general acid effects. The enhancement factor is ~  
205 1.5 at ionic strengths of 2 to 6.5 molal. This is relatively good agreement given that we  
206 estimate that the uncertainties in our rates is on the order of a factor of two, when  
207 considering individual uncertainties in the kinetics plots, sulfate quantification, and  
208 determination of AWC. The enhancement factor increases to  $33 \pm 4$  to  $51 \pm 9$  at ionic  
209 strengths of 14 molal. For comparison, the modeled enhancement factors for ionic  
210 strength of 0 to 5 molal solutions were calculated using parameters obtained in bulk  
211 solutions (34, 36, 37); see *SI Appendix*, section S2 for details. The model results were  
212 also extrapolated to ionic strength of up to 43 molal using the functional formula  
213 derived in the past using much more dilute solutions. The high enhancement factors at  
214 high ionic strength ranges, which are within those (13 through 43 molal) of urban  
215 pollution episodes (6), indicate that ionic strength effects on the aqueous oxidation of  
216 SO<sub>2</sub> by H<sub>2</sub>O<sub>2</sub> in aerosol particles should be considered in air quality models.

## 217 **Discussion**

218 By directly investigating the kinetics of aqueous oxidation of dissolved SO<sub>2</sub> by H<sub>2</sub>O<sub>2</sub> in  
219 aerosol particles under well-controlled experimental conditions, we have examined  
220 solute strength effects on the sulfate formation rate. The major result is that there is an  
221 enhancement in the kinetics above the rate that prevails in dilute solutions typical of  
222 cloudwater. We believe that the enhancement effect is due to ionic strength effects and  
223 general acid catalysis arising from the malonic acid buffer, as well as potentially to  
224 salting-in of the reactants.

225 Overall, the rate of this multiphase reaction is driven by the volume of liquid water  
226 present. When clouds are present, the liquid water content is many orders of magnitude  
227 higher than in aerosol particles, even for highly polluted conditions. Under such  
228 conditions, SO<sub>2</sub> oxidation will proceed preferentially in the cloud droplets. Similarly,  
229 this aerosol multiphase reaction is unlikely to be important in clean conditions when  
230 the aerosol liquid water content is very low. However, it has the potential to have  
231 atmospheric importance under polluted conditions.

232 To illustrate, Fig. 3 shows the steady-state sulfate formation rates by aqueous SO<sub>2</sub>  
233 oxidation through different reaction pathways in aerosol particles following the  
234 approach of Cheng et al. (6). According to their work and that of Zheng et al. (17), a  
235 missing sulfate source of anywhere from 0.3 to 5 μg m<sup>-3</sup> ·h<sup>-1</sup>, depending on the PM<sub>2.5</sub>  
236 level, is needed to explain the sulfate formation during haze episodes in Beijing in  
237 January 2013 (see figure 3 of ref. (6)). Based on this model (6), the NO<sub>2</sub> reaction  
238 pathway may be the missing sulfate source if the aerosol pH is above 6 while the TMI

239 reaction pathway will dominate sulfate formation if the aerosol pH is below 4.5.  
240 However, ionic strength effects on sulfate formation rate were not taken into account in  
241 the model of Cheng et al. (6).

242 By incorporating the enhancement factors of sulfate formation rate at high ionic  
243 strength (around 14 molal) into the model, the sulfate formation rate for the H<sub>2</sub>O<sub>2</sub>  
244 reaction pathway increases from ~0.07 μg m<sup>-3</sup> h<sup>-1</sup> to 2.3–3.6 μg m<sup>-3</sup> h<sup>-1</sup>. Recently, Ye et  
245 al. (18) have reported that H<sub>2</sub>O<sub>2</sub> concentrations during Beijing haze events were more  
246 than 1 order of magnitude higher than the value (0.01 ppb) assumed in the model of  
247 Cheng et al. (6). By updating the model with a H<sub>2</sub>O<sub>2</sub> concentration of 0.1 ppb, the sulfate  
248 formation rate for the H<sub>2</sub>O<sub>2</sub> reaction pathway increases to 23.2–36.0 μg m<sup>-3</sup> h<sup>-1</sup>.

249 We also investigated the effects of ionic strength on the aqueous TMI-catalyzed  
250 oxidation rate of dissolved SO<sub>2</sub> by O<sub>2</sub> in aerosol particles. We find that the sulfate  
251 formation rate is slower by a factor of ~ 85 at an ionic strength of 2.8 molal compared  
252 to that in dilute solution; see *SI Appendix*, section S4 for details. By taking the impact  
253 of ionic strength on the TMI reaction pathway into consideration, we find that the  
254 sulfate formation rate for the H<sub>2</sub>O<sub>2</sub> reaction pathway is larger than all other reaction  
255 pathways for aerosol pH levels up to 6.2. This pH value exceeds the reported aerosol  
256 pH values (4 to 5) during severe haze episodes in northern China (see ref. (38) and  
257 references therein). As the sulfate formation rate for the H<sub>2</sub>O<sub>2</sub> reaction pathway is  
258 independent of aerosol pH for pH levels above 2, the sulfate formation can be  
259 maintained at a high rate even when the aerosol particles become more acidic.

260 These results draw attention to the rates of formation of H<sub>2</sub>O<sub>2</sub> that are required to  
261 sustain the aerosol-mediated SO<sub>2</sub> oxidation. To our knowledge, the mixing ratio of H<sub>2</sub>O<sub>2</sub>  
262 was not measured for the January 2013 severe haze conditions presented in Fig. 3 (6).  
263 Modeling of H<sub>2</sub>O<sub>2</sub> measurements in Beijing during March 2016 was consistent with a  
264 production rate of 0.5 ppbv d<sup>-1</sup>, with more H<sub>2</sub>O<sub>2</sub> formed via alkene ozonolysis than via  
265 the HO<sub>2</sub> self-reaction (39). Assuming that H<sub>2</sub>O<sub>2</sub> production limits the rate of sulfate  
266 formation, this H<sub>2</sub>O<sub>2</sub> production rate can account for up to 20% of the missing sulfate  
267 source in the January 2013 case and roughly 1/3 of the sulfate formed in January and  
268 February 2015 reported by Wang et al. (4). These estimates are highly uncertain,  
269 however, being dependent on the photochemical conditions that give rise to H<sub>2</sub>O<sub>2</sub>.

270 Our results highlight the important role of high solute strength in sulfate formation  
271 in deliquesced aerosol particles. Specifically, there is the need for an improved  
272 understanding on the role of ionic strength effects on atmospheric aqueous multiphase  
273 chemistry, as pointed out by a recent review of the field of laboratory atmospheric  
274 chemistry (40). As well, the importance of general acid catalysis needs to be considered  
275 not only for the many organic acids that are present under pollution conditions but  
276 potentially also for ammonium ions which are present in aerosol particles at very high  
277 concentrations. While such rate enhancements may arise under any conditions with a  
278 high aerosol liquid water content and solute concentrations, for the specific Chinese  
279 haze situation it is particularly important to assess these effects on the reactions of  
280 organic hydroperoxides, NO<sub>2</sub> and O<sub>3</sub> that also oxidize dissolved SO<sub>2</sub> (41), over a range

281 of aerosol pH. This will permit a more comprehensive assessment of the overall rate of  
282 multiphase sulfur oxidation under cloud-free conditions.

## 283 **Methods**

284 **Aerosol particle generation.** Our goal was to investigate the kinetics of aqueous oxidation of  
285 dissolved SO<sub>2</sub> by H<sub>2</sub>O<sub>2</sub> in high solute strength aerosol particles, achieved by generating pH-buffered  
286 polydisperse deliquesced aerosol particles using a constant output atomizer (TSI 3076) from the  
287 following solutions: A) a mixture of NaCl/malonic acid/sodium bimalonate (10/0.5/0.5 mM,  
288 3/0.5/0.5 mM, and 1/0.5/0.5 mM); B) NaCl/sodium bimalonate/sodium malonate (10/0.5/0.5 mM,  
289 3/0.5/0.5 mM, and 1/0.5/0.5 mM); C) NaNO<sub>3</sub>/malonic acid/sodium bimalonate (10/0.5/0.5 mM); D)  
290 NaNO<sub>3</sub>/sodium bimalonate/sodium malonate (10/0.5/0.5 mM); E) malonic acid/sodium bimalonate  
291 (5/5 mM); and F) sodium bimalonate/sodium malonate (5/5 mM). A small fraction of the humidified  
292 aerosol flow from the atomizer was mixed with a humidified N<sub>2</sub> flow and entered the kinetics flow  
293 tube, resulting in high relative humidity (73-90%) to ensure that the seed aerosol particles are  
294 deliquesced and maintain enough aerosol liquid water to enable aqueous oxidation. Additional  
295 details on generating seed aerosols are provided in *SI Appendix*, section S1.

296 **Flow tube experiments.** A schematic of the full experimental setup is shown in *SI Appendix*, Fig.  
297 S1. All experiments were conducted in a vertically oriented pyrex flow tube (length 95 cm, inner  
298 diameter 6.2 cm) at a total flow rate of 1500 sccm at laminar flow conditions (Reynolds Number  
299 ~34), in which constant flows of SO<sub>2</sub> (10, 25, or 50 sccm carrier gas), gas phase H<sub>2</sub>O<sub>2</sub> (10 or 20  
300 sccm carrier gas), and aerosol particles were well mixed. Gas-phase H<sub>2</sub>O<sub>2</sub> was generated by  
301 bubbling N<sub>2</sub> through an aqueous H<sub>2</sub>O<sub>2</sub> solution (15% or 30% by weight) and injected into the central  
302 portion of the humidified aerosol flow by a movable stainless-steel injector tube that is inserted

303 axially down the center of the flow tube, enabling variable reaction time in a step-wise manner. In  
304 each experiment, the aerosol particles were first characterized in the presence of SO<sub>2</sub> and absence  
305 of H<sub>2</sub>O<sub>2</sub> to quantify the background sulfate concentrations. Then the aqueous oxidation of SO<sub>2</sub> was  
306 initiated by introducing H<sub>2</sub>O<sub>2</sub>. Before entering the instruments for aerosol composition  
307 characterization, the aerosol particles alternatively passed through or bypassed a diffusion dryer to  
308 remove or sustain aerosol water. Experiments with different initial concentrations of SO<sub>2</sub> and H<sub>2</sub>O<sub>2</sub>  
309 as well as seed aerosol types were carried out at 21 to 25 °C and high relative humidity (73 to 90%).  
310 All reported sulfate concentrations were corrected for the background sulfate concentrations. Details  
311 of the experimental setup and procedure are provided in *SI Appendix*, section S1.

312 **Measurements.** Both compact and high-resolution time-of-flight aerosol mass spectrometers (33)  
313 (C-ToF-AMS and HR-ToF-AMS; Aerodyne Research) were deployed to characterize the  
314 concentration and chemical composition of aerosol particles. Prior to the experiments, the ionization  
315 efficiency (IE) of the AMS was calibrated using 300 nm ammonium nitrate particles. The sulfate  
316 was measured in the form of sodium sulfate because the sulfuric acid that was generated in the  
317 reaction was buffered by sodium bimalonate or sodium malonate. Therefore, a sulfate fragmentation  
318 table without water fragments was used (42). The relative ionization efficiency (RIE) of sulfate was  
319 determined for sodium sulfate to be 0.12 and 0.24 for the C-ToF-AMS and HR-ToF-AMS,  
320 respectively, and applied to the quantification of sulfate concentrations. The difference between  
321 sulfate concentrations simultaneously measured by these two AMS was within 10%. A collection  
322 efficiency of unity was assumed for the deliquesced aerosol particles (43). An SMPS (TSI) was used  
323 to measure the particle size distributions. Gas-phase monitors were used to measure SO<sub>2</sub> (Model  
324 43i; Thermo) and H<sub>2</sub>O<sub>2</sub> (PI2114; PICARRO) gas-phase mixing ratios. Details regarding the



325 comprehensive calibration and operation of the AMS are described in *SI Appendix*, section S1.

326 **Calculation of aerosol pH, sulfate molality and ionic strength.** The relative concentrations of the  
327 solutes in the aerosol particles were assumed to be the same as in the solution in the atomizer with  
328 no fractionation, resulting in aerosol particles pH buffered at 2.3 to 4.8, estimated using the E-AIM  
329 model (29, 30) and a Pitzer activity coefficient model (31). Here, pH is defined as the negative  
330 logarithm with base 10 of the molality-based  $H^+$  activity.

$$331 \quad \text{pH} = -\log_{10}(\gamma_{H^+} m_{H^+}) \quad (3)$$

332 where  $\gamma_{H^+}$  and  $m_{H^+}$  represent the molality-based activity coefficient and molality of  $H^+$ , respectively.

333 For the mixture of NaCl and organic buffer, the molalities of all ions and neutral species were first  
334 estimated using the E-AIM model which were then input to the Pitzer model to calculate the  
335 equilibrium speciation to determine the aerosol pH. Given that the Pitzer model is developed for  
336 systems with NaCl and relatively low concentrations of malonic acid species, the aerosol pH for  
337 other systems (i.e. with no NaCl present) was estimated using the E-AIM model only. For  
338 comparison, the Pitzer model gave approximately 0.4 unit lower pH values for the mixture of  
339 NaCl/malonic acid/sodium bimalonate and 1 unit higher pH values for the mixture of NaCl/sodium  
340 bimalonate/sodium malonate compared to the E-AIM model results (*SI Appendix*, Fig. S9). We note  
341 that 1 unit difference of pH will not impact our conclusions given that the enhancement factors for  
342 the proton-catalyzed sulfate formation rate at the highest ionic strength (~14 molal) are pH-  
343 independent (*SI Appendix*, section S5 and Fig. S10A).

344 The sulfate molality was calculated from the sulfate concentrations ( $\mu\text{g m}^{-3}$ ) and aerosol liquid  
345 water volume, which was determined as the volume difference between the deliquesced and  
346 effloresced aerosol particles bypassing or passing through the diffusion dryer. We note that there is

347 an inherent assumption that the molar volumes of the solutes are the same in the dry and wet cases.  
348 For comparison, the aerosol liquid water volume was also determined by multiplying the total  
349 measured aerosol volume by the ratio of the aerosol liquid water volume to the total aerosol volume  
350 estimated using the E-AIM model, resulting in an enhancement factor of  $19 \pm 3$  to  $30 \pm 5$  for the  
351 proton-catalyzed sulfate formation rate at ionic strengths of 14 molal (*SI Appendix*, section S5 and  
352 Fig. S10B). This does not impact our conclusion that the oxidation of  $\text{SO}_2$  by  $\text{H}_2\text{O}_2$  in aerosol  
353 particles can contribute to the missing sulfate source during severe haze episodes.

354 The ionic strength was calculated via the following equation:

$$355 \quad I = \frac{1}{2} \sum m_i z_i^2 \quad (4)$$

356 where  $m_i$  and  $z_i$  represent the molality of each ion and its corresponding charge, respectively. The  
357 ionic strength was estimated using the E-AIM model (29) assuming the relative composition of the  
358 aerosol particles to be the same as in the solution (*SI Appendix*, Table S1).

359 **Data availability.** All data are available from the corresponding author on request.

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## 363 **Author contributions**

364 T.L. and J.A. designed the research project; T.L. performed the research; T.L., S.C., and J.A.  
365 analyzed data and T.L. and J.A. wrote the paper.

## 366 **Competing interests**

367 The authors declare no competing interests.

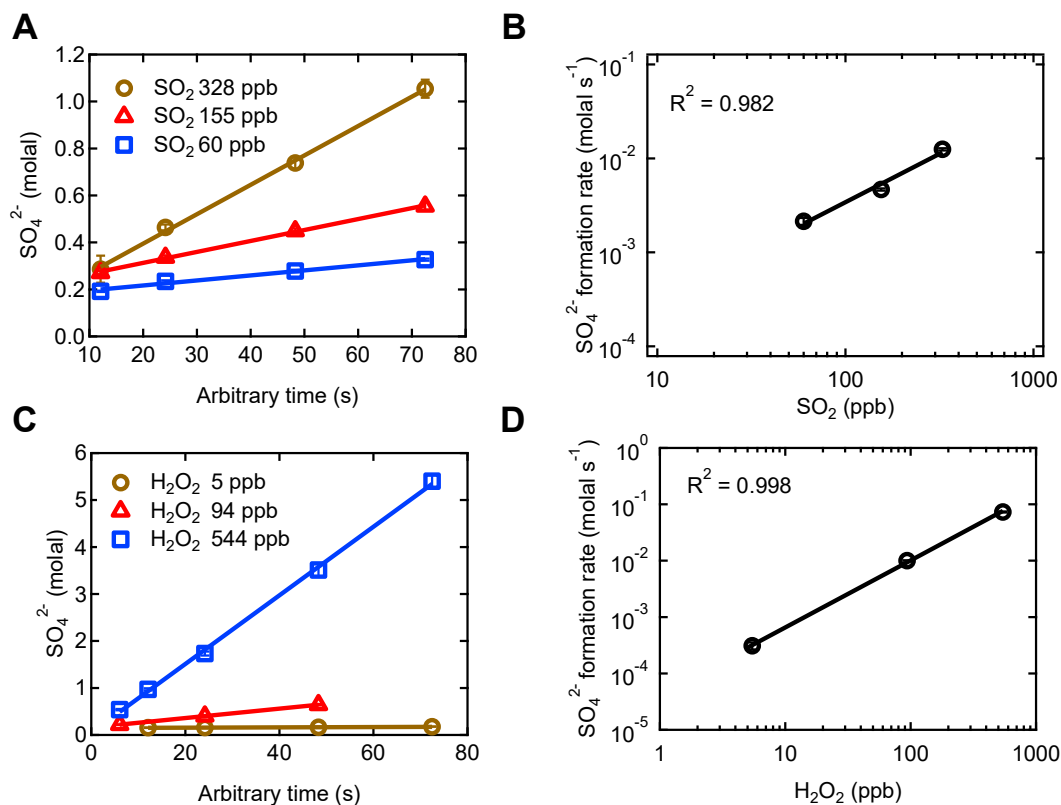
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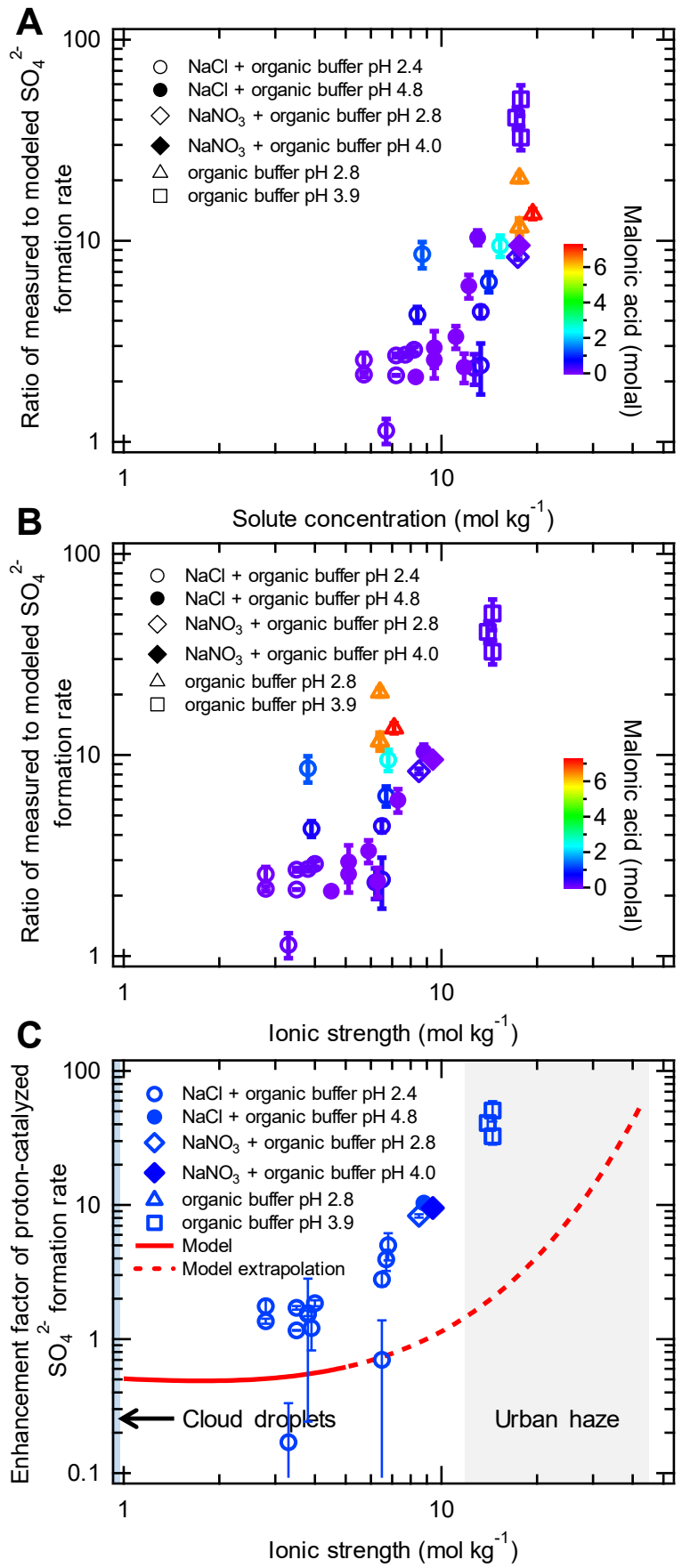
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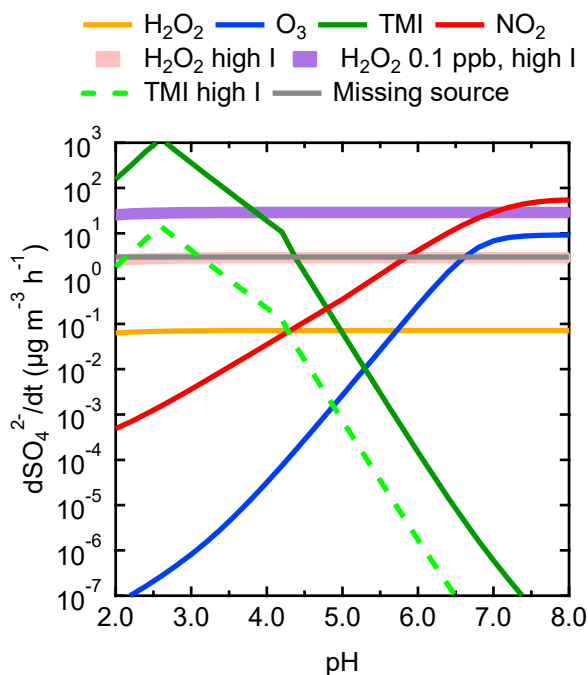
471

472 **Fig. 1. Kinetics of aqueous oxidation of SO<sub>2</sub> by H<sub>2</sub>O<sub>2</sub> in a mixture of NaCl and**  
 473 **malonic acid buffer aerosol particles (molar ratio of 20:1:1) at pH 2.5. (A)** Sulfate  
 474 concentrations (units of molal) as a function of reaction time for initial SO<sub>2</sub>  
 475 concentrations of 60, 155, and 328 ppb, for a H<sub>2</sub>O<sub>2</sub> mixing ratio of 94 ppb (*SI Appendix*,  
 476 Table S1, Exp 7 through 9). **(B)** Dependence of sulfate formation rate (molal s<sup>-1</sup>) on  
 477 initial SO<sub>2</sub> concentrations, for a H<sub>2</sub>O<sub>2</sub> mixing ratio of 94 ppb. **(C)** Sulfate concentrations  
 478 as a function of reaction time for initial H<sub>2</sub>O<sub>2</sub> concentrations of 5, 94, and 544 ppb, for  
 479 a SO<sub>2</sub> mixing ratio of 340 ppb (*SI Appendix*, Table S1, Exp 10 through 12). **(D)**  
 480 Dependence of sulfate formation rate on initial H<sub>2</sub>O<sub>2</sub> concentrations, for an SO<sub>2</sub> mixing  
 481 ratio of 340 ppb. All error bars represent 1 SD; they are generally smaller than the  
 482 symbol size.



484 **Fig. 2. Ratio of the measured to modeled sulfate formation rates and enhancement**  
485 **factor of proton-catalyzed sulfate formation rate. (A)** Dependence of the ratio of the  
486 measured to modeled sulfate formation rate on total solute concentration and malonic  
487 acid concentrations for a mixture of NaCl and organic buffer aerosol particles (circle  
488 symbols), NaNO<sub>3</sub> and organic buffer aerosol particles (rhombus symbols) and organic  
489 buffer aerosol particles at aerosol pH of 2.8 (triangle symbols) and 3.9 (square symbols).  
490 **(B)** Same as Fig. 2A but plotted against ionic strength. See *SI Appendix*, Fig. S4 for an  
491 equivalent plot where RH is indicated. **(C)** Dependence of the enhancement factor of  
492 proton-catalyzed sulfate formation rate on ionic strength. The calculated proton-  
493 catalyzed sulfate formation rates for organic buffer aerosol particles at aerosol pH of  
494 2.8 and most data for the NaCl and organic buffer aerosol particles at aerosol pH of 4.8  
495 were negative and so are not plotted. The model results (solid line) describe the effects  
496 of ionic strength on the proton-catalyzed sulfate formation rate measured in bulk  
497 solutions (34, 36, 37). The model extrapolation is shown as a dashed red line. Light  
498 blue- and gray-shaded areas indicate characteristic ionic strength ranges for cloud  
499 droplets and aerosol particles during urban haze episodes, respectively. All error bars  
500 are derived from the uncertainties of the measured sulfate formation rates.





501

502 **Fig. 3. Sulfate formation rates by aqueous SO<sub>2</sub> oxidation in aerosol particles under**

503 **severe haze conditions.** Colored solid lines represent sulfate formation rates calculated

504 for different reaction pathways at conditions of: [SO<sub>2</sub> (g)] = 40 ppb, [H<sub>2</sub>O<sub>2</sub> (g)] = 0.01

505 ppb, [NO<sub>2</sub> (g)] = 66 ppb, [O<sub>3</sub> (g)] = 1 ppb, pH-dependent concentrations of [Fe (III)]

506 and [Mn (II)] (6), AWC = 300 µg m<sup>-3</sup>, aerosol droplet radius R<sub>p</sub> = 0.15 µm, T = 271 K,

507 and I = 0 molal. The total soluble Fe and Mn are 18 and 42 ng m<sup>-3</sup>, respectively. The Fe

508 (III) concentration may be limited by the solubility of Fe(OH)<sub>3</sub> in which case the

509 saturated concentration of Fe(III) was estimated from the solubility product equilibrium

510 constant of Fe(OH)<sub>3</sub> (K<sub>sp</sub> = 2.6 × 10<sup>-38</sup>) (44). The pink-shaded area represents the sulfate

511 formation rate for H<sub>2</sub>O<sub>2</sub> reaction pathway accounting for the rate enhancement at I = 14

512 molal (and [H<sub>2</sub>O<sub>2</sub> (g)] = 0.01 ppb). The purple-shaded area represents the sulfate

513 formation rate for H<sub>2</sub>O<sub>2</sub> reaction pathway at I = 14 molal and a higher concentration of

514 [H<sub>2</sub>O<sub>2</sub> (g)] = 0.1 ppb as observed during Beijing haze events (18). The dashed green

515 line indicates the sulfate formation rate for TMI reaction pathway accounting for the

516 impact of ionic strength. The grey line shows a missing sulfate formation rate of 3 µg

517 m<sup>-3</sup> · h<sup>-1</sup> as a reference. This figure is constructed using the characteristic data conditions

518 from ref. 6 along with additional components related to our analysis.